On polynomial predictions for river flows^{*}

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December 18, 2023

Abstract

This paper deals with the prediction of river levels by means of polynomial regression models using only elevation data and inflow forecasts. Different models for this purpose are examined and a new approach based on the concept of virtual stations is presented. Detailed numerical experiments show that this proposal may be useful as a tool for making predictions when the physical characteristics of the river are uncertain.

Key words: Flow predictions, natural rivers, Saint-Venant equations, parameter estimation.

1 Introduction

River flow modelling is an important tool for analysing and predicting dam failures and their consequences. The main mathematical procedure for this task is based on the solution of partial differential equations (PDE). The equations of Saint Venant [20] are the best known equations for this purpose. Their numerical solution requires initial and boundary conditions in terms of river wetted cross-sections and flow-rates. In addition, geometric descriptions of the cross sections and bed elevations are required. Finally, Manning roughness coefficients, which may be spatially and temporally dependent, must be determined. See [1, 2, 3, 5, 4, 6, 7, 8, 11, 12, 13, 14, 17, 18, 19, 20, 21].

Typically, partial observations of river surface elevations at different spatial and temporal coordinates are available. These observations make it possible the estimation of the unknown characteristics of the river, which are necessary for the numerical integration of the partial differential equations. The resulting PDE-constrained parameter estimation problem can be difficult to solve, requires integration of the PDE's for different instances, and is subject to instability and lack of reliability of results. However, this problem has been the subject of valuable research over many years. See [1, 2, 3, 5, 6, 7, 8, 11, 14, 15, 17, 18].

The PDE approach obtains predictions by means of the estimation of unknown physical characteristics and associated PDE integration. Moreover, the estimation of unknown physical characteristics is based on fitting the direct solution of the PDE's to available observations. This suggests the possibility of obtaining river predictions directly from available data without the need to estimate the physical characteristics of the river. The obvious drawback of this approach relies on the fact that

^{*}This work was supported by FAPESP (grants 2013/07375-0, 2018/24293-0, and 2022/05803-3) and CNPq (grants 302538/2019-4 and 302682/2019-8).

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we do not have reliable physical models that directly link observations to predictions. For this reason we believe that data-based predictions should generally be considered in conjunction with PDE predictions, although the specific form of this relation is highly problem-dependent [6].

Reliable data-based approaches should start with a reliable identification of cause-effect relationships. For example, in the case of river flow phenomena, a high correlation may be found between upstream discharge and downstream elevations. Obviously, upstream discharges are the cause of downstream elevations and not the other way round. If a cause-effect relationship is established, the next step could be to propose an appropriate form of dependence relationship, the specific form of which should be based on previous data analysis.

Let us consider an example that is well suited to introduce and motivate the rest of this paper. It has been widely observed that water elevation at an arbitrary fixed station of a natural river is a smooth function of the upstream (inlet) flow-rate. See [12] and [2, Fig.12b]. In Figure 1, we consider data for the Fork River published in [9]. Figure 1a shows observations of the elevation z corresponding to the section x = 751 m, together with linear, quadratic and cubic polynomials representing elevation as a function of the inflow rate Q_{\min} (in m^3/s). The polynomials were fitted using simple least squares. Figure 1b shows the same information but related to the section x = 3256 m. The observations are taken every 12 hours starting at zero hours on day 3. The polynomial coefficients and the corresponding root mean square deviation (RMSD) are given in Table 1.

Station	Polynomial	RMSD	c_0	c_1	c_2	c_3
В	linear	8.69579603E-02	7.35113673	3.75568519E-02		
21	quadratic	2.69668513E-02	7.08338033	8.19336547 E-02	-1.36086954E-03	
-1	cubic	2.42162234E-02	7.01642805	9.97412870 E-02	-2.60953038E-03	2.47226020E-05
Ш	linear	6.02123240E-02	5.44084782	3.91356263E-02		
256	quadratic	3.13462816E-02	5.28175904	6.62052107 E-02	-8.39381211E-04	
32	cubic	3.07813747E-02	5.24970397	7.49278445 E-02	-1.45802975E-03	1.23271897E-05

Table 1: Fork river: fitting polynomials, their coefficients, and the corresponding RMSD (in meters).

It is interesting to fit the data of, say, the first 10 days and observe if the approximating curves fit well the data for the remaining days. Figure 2 and Table 2 show the results. Throughout this paper surface elevations and the corresponding RMSD errors are expressed in meters. So, for example, the testing error of the cubic polynomial for x = 751 m meters is 4.80 cm according to Table 2. This error is quite small for practical prediction purposes regarding a real river.

Station	Polymomial	RM	ISD	0-	0.	0-	0-
Station	1 Orynolinai	training	testing		C1	C2	63
В	linear	4.36362260E-02	2.24297179E-01	7.66210301	2.34403945E-02		
51	quadratic	1.97586581E-02	1.11828660E-01	7.33545084	5.90878962 E-02	-8.67373515E-04	
	cubic	1.06148710E-02	4.79787043E-02	6.94435497	1.24958948E-01	-4.23241309E-03	5.33788086E-05
н	linear	4.41380828E-02	1.67298488E-01	5.67319985	2.89537322E-02		
556	quadratic	3.43615061E-02	8.69890046E-02	5.44061014	5.43362118E-02	-6.17605427E-04	
	cubic	2.73302025E-02	7.05794611E-02	4.95184090	1.36658084E-01	-4.82303941E-03	6.67097817 E-05

Table 2: Fork river: Fitting polynomials, their coefficients, and the corresponding RMSD. In this case, observations of the first 10 days were used as training data to fit the polynomials. The remaining observations (20 or 21 days for Sections x = 751 m and Section x = 3256 m, respectively) were not used in the fitting (training) phase and, then, were used to test the predictions provided by the fitted polynomials.



Figure 1: Fork river: Observed elevations at a given station and their approximation as a (linear, quadratic and cubic) fitting polynomial of the inlet discharge.

These results suggest that, for predicting elevations at a fixed station x in "future days" under suitable forecast on the inlet discharge, it is enough to fit the curve of the surface elevation z(x,t)versus $Q_{\min}(t)$ using available data at station x, with the reasonable belief that, in the next days, this



Figure 2: Fork river: Observed elevations at a given station and their approximation as a (linear, quadratic and cubic) fitting polynomial of the inlet discharge. In this case, observations of the first 10 days were used as training data to fit the polynomials. The remaining observations (20 or 21 days for Sections x = 751 m and Section x = 3256 m, respectively) were not used in the fitting (training) phase and, then, were used to test the predictions provided by the fitted polynomials.

curve will provide reasonable elevation estimates, provided that inlet discharge forecasts are reliable. In fact, this should be the case if one has data for a suitable number of days before "today" and for all the relevant stations along the river. Unfortunately both situations are unlike to occur. Usually, one needs previsions for the future employing a possibly moderate number of past data at a possibly moderate number of stations x.

For example, according to Table 1, for x = 751 m, the best third-order polynomial that represents z(x,t) as a function of $Q_{\min}(t)$ is given by

$$z(751,t) \approx 7.02 + 9.97 \times 10^{-2} Q_{min}(t) - 2.61 \times 10^{-3} Q_{min}(t)^2 + 2.47 \times 10^{-5} Q_{min}(t)^3, \tag{1}$$

while, for x = 3256 m, the best third-order polynomial that represents z(x, t) as a function of $Q_{\min}(t)$ is given by

$$z(3256,t) \approx 5.25 + 7.49 \times 10^{-2} Q_{min}(t) - 1.46 \times 10^{-3} Q_{min}(t)^2 + 1.23 \times 10^{-5} Q_{min}(t)^3.$$
(2)

However, if $x \notin \{751, 3256\}$, we do not know, for example, which is the best third-order polynomial that fits the elevations z(x, t) at Section x = 555 m as a function of $Q_{\min}(t)$. This question is addressed in the present paper.

We will start from the empirical observation that, in real rivers, inlet discharge is the dominant cause of river elevations at different stations. This fact supports the idea that, given a spatial position x, the elevation z(x,t) can be well approximated by a low-order polynomial $P(Q_{\min}(t))$. We will see that third-order polynomials are the more appropriate for this purpose. In order to recover elevations at stations x that are not represented in the data we analyse the employment of two-dimensional polynomials in the variables x and $Q_{\min}(t)$. However, the need to preserve the accuracy of the onedimensional fits leads us to propose a different strategy based on the concept of "virtual stations". This paper proposes an algorithm for selecting suitable virtual stations and demonstrates its reliability through detailed numerical experiments.

This research is conducted within CRIAB, a Latin-American academic group that involves collaborators of several countries. The group is dedicated to analyzing, comprehending and mitigating dam-breaking and related accidents. River modelling is one of the techniques that must be mastered in the broader landscape of modelling embankments and basins. Optimization regression techniques are among the tools used for this purpose.

This paper is organized as follows. Section 2 analyses the compatibility of one-dimensional regression with two-variable polynomial fitting. Section 3 introduces the method of virtual stations and describes the algorithm that will be used in the experiments. Section 4 describe the generation of synthetic data. Numerical experiments are reported in Section 5, while conclusions and future research directions are presented in Section 6.

Notation. #A will denote the number of elements of the set A. If A and B are sets, $A \setminus B$ denotes the set of elements of A that do not belong to B.

2 Two-variable polynomial fitting

Consider an arbitrary one-dimensional flow where the spatial (length) coordinate x goes from x_{\min} to x_{\max} . The surface elevation for space coordinate x and time coordinate t will be denoted z(x,t). Assume that at p different stations $x_1, \ldots, x_p \in [x_{\min}, x_{\max}]$ we have observations of surface elevations at different times. The inlet discharge (flow-rate at $x = x_{\min}$) at time $t \in [t_{\min}, t_{\max}]$ is denoted $Q_{\min}(t)$. For simplicity, if confusion is not possible, we omit the dependence of t in this notation (denoting $Q_{\min} = Q_{\min}(t)$). Assume that, at each station x_j , we fit a polynomial $P_j(Q_{\min})$ with degree q, in the least-squares sense, in order to minimize the deviations with respect to measured elevations.

We may consider the model

$$z(x,t) \approx W_1(x)P_1(Q_{\min}(t)) + \dots + W_p(x)P_p(Q_{\min}(t)),$$
 (3)

where, for all j = 1, ..., p, $W_j(x)$ is a polynomial with degree p-1 such that $W_j(x_j) = 1$ and $W_j(x_\ell) = 0$ if $\ell \neq j$. Namely,

$$W_{j}(x) = \frac{\prod_{i \neq j} (x - x_{j})}{\prod_{i \neq j} (x_{i} - x_{j})}.$$
(4)

The right-hand side of (3) is a sum of p(q+1) monomials of the form $\gamma_{i,j}x^iQ^j_{\min}$ for $i = 0, 1, \ldots, p-1$ and $j = 0, 1, \ldots, q$.

This suggests the model

$$z(x,t) \approx \sum_{i=0}^{s} \sum_{j=0}^{q} \gamma_{i,j} x^{i} Q_{\min}(t)^{j}.$$
(5)

In (5), we postulate that the elevation at each point (x, t) is a two-variable polynomial with variables x and $Q_{\min}(t)$, with degree s in the variable x and degree q in the variable Q_{\min} . Note that in (3) we have that s = p - 1.

The model (5) induces a linear least-squares problem, in which the coefficients $\gamma_{i,j}$ are the unknowns and observations are available at different stations and times. We wonder whether, if observations are given at a finite number of stations $x_1, \ldots x_p$, the solution of the least-squares problem comes from addressing p separate least squares problems, one corresponding to each station. In this case, we could compute the best polynomial of degree q with respect to measurements at the considered station and the predicted values at arbitrary points (x, t) would come from interpolation according to (3) and (4).

The following theorem gives an answer to this question.

Theorem 2.1 Assume that elevations $z_{k,\ell}$ are given at p stations x_k , k = 1, ..., p, and time instants t_ℓ , $\ell = 1, ..., r_k$. Assume, moreover, that for each observed $z_{k,\ell}$ the inlet flow $Q_{\min}(t_\ell)$ (in short Q_ℓ) is known. Consider the linear least-squares problems

Minimize
$$\sum_{k=1}^{p} \sum_{\ell=1}^{r_k} \left[\sum_{j=0}^{q} \sum_{i=0}^{s} \gamma_{i,j} x_k^i Q_\ell^j - z_{k,\ell} \right]^2$$
 (6)

and

Minimize
$$\sum_{k=1}^{p} \sum_{\ell=1}^{r_k} \left[\sum_{j=0}^{q} \beta_{k,j} Q_{\ell}^j - z_{k,\ell} \right]^2$$
. (7)

Then, the objective function value at the solution of (7) is less than or equal to the objective function value at the solution of (6). Moreover, if $s \ge p - 1$ both objective functions are identical at respective solutions.

Proof: Problem (6) is equivalent to

$$\text{Minimize} \sum_{k=1}^{p} \sum_{\ell=1}^{r_k} \left[\sum_{j=0}^{q} \beta_{k,j} Q_{\ell}^j - z_{k,\ell} \right]^2 \tag{8}$$

subject to

$$\beta_{k,j} = \sum_{i=0}^{s} \gamma_{i,j} x_k^i \text{ for all } k = 1, \dots, p, j = 0, 1, \dots, q.$$
(9)

Therefore, problem (6) is equivalent to problem (7) with the additional constraints (9). So, the feasible region of (7) contains the feasible region of (8,9). This implies that the objective function of (7) at its solution is smaller than or equal to the objective function of (8,9) at its solution. Both objective function values are identical if the feasible region of (7) is the same as the feasible region of (8,9), that is, if for all $\beta_{k,j} \in \mathbb{R}$ there exist $\gamma_{i,j}$ such that the identity (9) holds. This would mean that the linear system (9) (with unknowns $\gamma_{i,j}$) and independent term given by $\beta_{k,j}$) is compatible.

By (9), for j = 0, 1, ..., q, we have

$$\beta_{1,j} = \gamma_{0,j} x_1^0 + \gamma_{1,j} x_1^1 + \dots + \gamma_{s,j} x_1^s, \tag{10}$$

$$\beta_{2,j} = \gamma_{0,j} x_2^0 + \gamma_{1,j} x_2^1 + \dots + \gamma_{s,j} x_2^s, \tag{11}$$

$$\beta_{p,j} = \gamma_{0,j} x_p^0 + \gamma_{1,j} x_2^1 + \dots + \gamma_{s,j} x_p^s.$$

$$\tag{12}$$

If s < p-1 the systems (10)–(12) are overdetermined and the solution set may be empty. In that case, the objective function value at the solution of (6) could be bigger than the objective function value at the solution of (7). If s = p-1, for each $j = 0, 1, \ldots, q$, the equations (10)–(12) define a $p \times p$ Vandermonde system. See [10, pp.203-207]. So, the q + 1 systems (10–(12) are compatible and the unknowns $\gamma_{0,j}, \ldots, \gamma_{p-1,j}$ are (uniquely) determined by the constraints (9). If s > p-1 the systems (10)–(12 are underdetermined and particular solutions come from completing the solutions of the case s = p-1 with $\gamma_{p,j} = \ldots, \gamma_s = 0$. Therefore, when $s \ge p-1$, the constraints (9) do not impose any constraint at all to the solution of (8). Thus, the problems (6) and (7) are equivalent when $s \ge p-1$. This completes the proof.

. . .

However, if observations $z_{obs}(x_k, t_k)$ are available at different times and stations (x_k, t_k) , $k \in K_{obs}$, we must rely directly on the least squares problems induced by (5). Namely,

Minimize
$$\sum_{k \in K_{\text{obs}}} \left[\sum_{i=0}^{s} \sum_{j=0}^{q} \gamma_{i,j} x_k^i Q_{\min}(t_k)^j - z_{\text{obs}}(x_k, t_k) \right]^2.$$
 (13)

Note that problems of the form (6) are of the form (13) but the reciprocal is not true. Observe, moreover, that the number of parameters γ_{ij} that are estimated when we use (13) is (s+1)(q+1), where s is the degree of the polynomial with respect to the variable x and q is the degree of the polynomial with respect to the variable x and q is the degree of the polynomial with respect to the variable x and q is the degree of the polynomial with respect to the variable x and q is the degree of the polynomial with respect to the variable x.

3 Method of virtual stations

Assume that we have p observation stations with spatial coordinates x_1, \ldots, x_p and that, for all $i = 1, \ldots, p$, N_i elevation observations are available for N_i different temporal coordinates. It is plausible that, as suggested in Section 1, and as will be confirmed by forthcoming experiments, the best model for the predicted elevations at any given station should come from a least-squares fitting of a suitable polynomial using the observed associated elevations. If the degree of each polynomial is q, the number of coefficients of this model is p(q + 1). It is disappointing that this number is, in

general, bigger than (s+1)(q+1), which is the number of coefficients associated with the two-variable polynomial model discussed in Section 2. Therefore, solving (13) does not lead to the likely optimal elevation prediction, given the data availability mentioned in this paragraph.

On the other hand, the procedure based on (13) seems to be suitable for the case where one has observations at different space-time positions, not necessarily concentrated at fixed stations. In this section we will assume that available elevation data $z_{obs}(x_k, t_k)$ are given at n_{dat} space-time points (x_k, t_k) for $k = 1, \ldots, n_{dat}$. We also assume that inlet discharge $Q_{\min}(t)$ is available whenever necessary.

We consider that $x_{\min} \leq \bar{x}_1 < \bar{x}_2 < \cdots < \bar{x}_{n_{\text{stat}}} \leq x_{\max}$. Each spatial position \bar{x}_j will be called "virtual station". The unknowns of our problem will be the coefficients $c_{0,j}, c_{1,j}, c_{2,j}, c_{3,j}$ for all $j = 1, \ldots n_{\text{stat}}$. Note that our fitting problem has $4n_{\text{stat}}$ unknowns. The objective function f will be a sum of squared errors, each error corresponding to an elevation observation. Namely,

$$f(c) = \sum_{k=1}^{n_{\text{dat}}} \left[z_{\text{cal}}(x_k, t_k, c) - z_{\text{obs}}(x_k, t_k) \right]^2,$$
(14)

where c is the vector of estimated coefficients c_{ij} stored columnwise and $z_{cal}(x_k, t_k, c)$ is the elevation computed by the model at the point (x_k, t_k) when the model coefficients are given by the vector c.

Let us describe how $z_{cal}(x_k, t_k, c)$ is computed. Given $k \in \{1, \ldots, n_{dat}\}$ we define $x_{left(k)}$ as the biggest \bar{x}_j such that $\bar{x}_j \leq x_k$ and we define $x_{right(k)}$ as the smallest \bar{x}_j such that $x_k < \bar{x}_j$, except in the cases that $x_k < \bar{x}_1$ or $x_k > \bar{x}_{n_{stat}}$. If $x_k < \bar{x}_1$ we define $x_{left(k)} = \bar{x}_1$ and $x_{right(k)} = \bar{x}_2$. If $x_k > \bar{x}_{n_{stat}}$ we define $x_{left(k)} = \bar{x}_1$ and $x_{right(k)} = \bar{x}_2$. If $x_k > \bar{x}_{n_{stat}}$ and $c_{0,right(k)}, c_{1,right(k)}, c_{2,right(k)}, c_{3,right(k)}$ will be the only coefficients that appear in the definition of $z_{cal}(x_k, t_k, c)$.

We define

$$w_{\text{left}(k)} = c_{0,\text{left}(k)} + c_{1,\text{left}(k)}Q_{\min}(t_k) + c_{2,\text{left}(k)}Q_{\min}(t_k)^2 + c_{3,\text{left}(k)}Q_{\min}(t_k)^3$$
(15)

and

$$w_{\text{right}(k)} = c_{0,\text{right}(k)} + c_{1,\text{right}(k)}Q_{\min}(t_k) + c_{2,\text{right}(k)}Q_{\min}(t_k)^2 + c_{3,\text{right}(k)}Q_{\min}(t_k)^3.$$
(16)

Finally,

$$z_{\rm cal}(x_k, t_k, c) = \frac{x_k - x_{\rm right(k)}}{x_{\rm left(k)} - x_{\rm right(k)}} w_{\rm left(k)} + \frac{x_k - x_{\rm left(k)}}{x_{\rm right(k)} - x_{\rm left(k)}} w_{\rm right(k)}.$$
(17)

According to (15), (16), and (17), $z_{cal}(x_k, t_k, c)$ depends linearly on the unknown coefficients c. Therefore, the minimization of (14) is a linear least-squares problem. This problem has n_{dat} equations and $4n_{stat}$ unknowns. Note that the number of virtual stations and their positions are arbitrary and should be chosen taken into account the coordinates of the available data.

3.1 Choosing virtual stations

The positions of the virtual stations $\bar{x}_1, \ldots, \bar{x}_{n_{\text{stat}}} \in [x_{\min}, x_{\max}]$ are "hyper-parameters" of the model presented in Section 3. The objective function in the model "with variable virtual stations" is given by (14) and each $z_{\text{cal}}(x_k, t_k, c)$ is defined by (17), but $x_{\text{right}(k)}$ and $x_{\text{left}(k)}$ are now variables of the problem that may change in order to obtain better values of the objective function. Therefore, a more precise definition of the objective function is

$$f(c,\bar{x}) = \sum_{k=1}^{n_{\text{dat}}} \left[z_{\text{cal}}(x_k, t_k, c) - z_{\text{obs}}(x_k, t_k) \right]^2,$$
(18)

where the coordinates of \bar{x} are $\bar{x}_1, \ldots, \bar{x}_{n_{\text{stat}}}$ and, for all $k = 1, \ldots, n_{\text{dat}}$,

$$z_{\rm cal}(x_k, t_k, \bar{x}, c) = \frac{x_k - x_{\rm right}(k)}{x_{\rm left}(k) - x_{\rm right}(k)} w_{\rm left}(k) + \frac{x_k - x_{\rm left}(k)}{x_{\rm right}(k) - x_{\rm left}(k)} w_{\rm right}(k).$$
(19)

Let us define now an algorithm that we effectively use for choosing the coordinates of stations $\bar{x}_1, \ldots, \bar{x}_{n_{\text{stat}}}$. Let us initialize the set \mathcal{O} in the following way:

$$\mathcal{O} = \{ x \in [x_{\min}, x_{\max}] \text{ such that there exists } k \in \{1, \dots, n_{dat}\} \text{ with } x = x_k \}.$$
(20)

Note that we could define

$$\mathcal{O} = \{x_1, \dots, x_{n_{\text{dat}}}\},\$$

but this definition should be ambiguous, inducing that the number of elements of \mathcal{O} is n_{dat} . This is not the case, because *x*-coordinates may be repeated in the set of observations. In fact, the number of elements of \mathcal{O} is less than or equal to n_{dat} . From now on, we will assume that the cardinality of \mathcal{O} is not smaller than 2. Therefore, one has at least two values of spatial coordinates *x* for which we have at least one observation. Note that the number of elements of \mathcal{O} is between 2 and n_{dat} and that this number may be strictly smaller than n_{dat} . The set of positions of the virtual stations will be called \mathcal{S} . It will be defined recursively in the following way:

Algorithm 3.1.1. Initialize $\mathcal{S} \leftarrow \emptyset$.

Step 1. If $\#S \ge n_{\text{stat}}$ or $\mathcal{O} = \emptyset$, stop.

Step 2. Compute a solution \hat{x} of the problem

$$\underset{x \in \mathcal{O}}{\text{Maximize min}} \left\{ \# \mathcal{L}(x), \# \mathcal{R}(x) \right\}$$
(21)

where

$$\mathcal{L}(x) := \{k \in \{1, \dots, n_{\text{dat}}\} \mid x_{\text{left}(k)} = x\} \text{ and } \mathcal{R}(x) := \{k \in \{1, \dots, n_{\text{dat}}\} \mid x_{\text{right}(k)} = x\}$$

Step 3. Update $\mathcal{S} \leftarrow \mathcal{S} \cup \{\hat{x}\}$ and $\mathcal{O} \leftarrow \mathcal{O} \setminus \{\hat{x}\}$ and go to Step 1.

At each iteration, the algorithm chooses the virtual station that maximizes the minimum number of available observations to determine each of the n_{stat} station cubic polynomial by means of leastsquare calculations. It is clear that, after a finite number of steps we have that the number of elements of S is n_{stat} or that \mathcal{O} is empty and the algorithm stops.

4 Generation of synthetic data

In order to evaluate the effectiveness of different regression models for river predictions, we need to rely on synthetic experiments. In our present research we decided to generate synthetic data by means of integration of the Saint-Venant equations [20], which are given by

$$\frac{\partial A}{\partial t} + \frac{\partial Q}{\partial x} = 0 \tag{22}$$

and

$$\frac{\partial Q}{\partial t} + \frac{\partial}{\partial x} \left(\frac{Q^2}{A}\right) + gA\frac{\partial z}{\partial x} + \frac{n_g^2 Q|Q|}{AR^{4/3}} = 0$$
(23)

for $x \in [x_{\min}, x_{\max}]$ and $t \in [t_{\min}, t_{\max}]$, where $h(x, t) = z(x, t) - z_b(x)$ is the depth of the river at (x, t), A(x, t) = h(x, t) w(x) is the cross wetted area at (x, t), P(x, t) = w(x) + 2h(x, t) is the wetted perimeter at (x, t), R(x, t) = A(x, t)/P(x, t) is the hydraulics radius at (x, t), V(x, t) = Q(x, t)/A(x, t) is the speed of the fluid at (x, t), and g is the acceleration of gravity taken as $9.81m/s^2$. Equation (22) describes mass conservation and equation (23) represents conservation of the linear momentum. The coefficient n_g is known as Manning roughness coefficient. It is unclear in which way this coefficient depends on x or t. On the one hand, the roughness coefficient depends on x due to the morphological differences of the river along its course. On the other hand, sediment deposition can also affect the roughness coefficients over time. In (23), n_g has units $m^{1/6}$.

The Saint-Venant equations were solved approximately by means of an explicit diffusive finitedifference method [13, 19] with the following specifications:

- $x_{\min} = 0$ and $x_{\max} = 3000$ (meters).
- $t_{\min} = 0$ and $t_{\max} = 29 + \frac{23}{24}$ (days) or, equivalently, 719 hours or 2,588,400 seconds.
- Initial conditions $z(x, t_{\min})$ given in Figure 3 and $Q(x, t_{\min}) = 3.9 \text{ m}^3/\text{s}$ for all $x \in [x_{\min}, x_{\max}]$.
- Boundary condition $Q(x_{\min}, t)$ given in Figure 4.
- Manning coefficient $n_g(x) = 0.078$ for all $x \in [x_{\min}, x_{\max}]$.
- Time step $\Delta t = 1$ second, spatial step $\Delta x = 30$, and diffusion coefficient 0.99.

Note that, according to the considered discretization, the finite difference method computes the values of z(x,t) and Q(x,t) at 101×2588401 points. We store only the values of z(x,t) and Q(x,t) for $x = 0, 30, 60, \ldots, 3000$ meters and for $t = 0, 1, 2, \ldots, 719$ hours. In other words, the "observed" elevations are given by a matrix of 101×720 positions. The level sets defined by this matrix is given in Figure 5.

5 Numerical Experiments

The data used in the numerical experiments are generated as described in Section 4. The employment of synthetic data allows us to test regression models in situations in which real data are not available.

5.1 Single-station one-dimensional models

In this short subsection, using synthetic data, we perform the same one-dimensional models experiment described in Section 1. In this case we use the stations defined by x = 720 m and x = 3000 m. We wish to verify whether the performance of the polynomial one-dimensional models for reproducing synthetic data is similar to the performance reported for real data in Section 1. Figures 6 and 7 and Tables 3 and 4 show the results. Clearly, in terms of quality of fitting and predictions, the performance of the polynomial models using synthetic data is similar to the one that has been reported in Section 1 for data of the real Fork River.

5.2 Experiments using observations on a mesh

In this subsection we consider as observations data between days t = 3 and t = 1010, every 12 hours, at 26 equally spaced stations between $x_{\min} = 0$ and $x_{\max} = 3000$ meters. The objective is, with these meshed data, to predict the elevation z(x, t) at 26 the equally spaced stations between $x_{\min} = 0$ and $x_{\max} = 3000$ meters and $t \in \{11, 12, \ldots 29\}$. We consider six different ways of prediction by combining



Figure 3: Initial condition for z used in the generation of synthetic data.



Figure 4: Q boundary condition used in the generation of synthetic data.

two types of polynomials (interpolating and least squares) and three possible degrees (linear, quadratic and cubic). Specifically, each of the six experiments consists of:



Figure 5: Synthetic Elevations.

Station	Polynomial	RMSD	c_0	c_1	c_2	c_3
В	linear	1.68217006E-02	7.36053069	3.42611821E-02		
201	quadratic	3.84570654E-03	7.30939110	4.29580343E-02	-2.70630819E-04	
	cubic	2.59511479E-03	7.29329906	4.73743343E-02	-5.86479099E-04	6.35214794 E-06
В	linear	4.19519692E-02	5.81624209	2.58508659E-02		
00	quadratic	1.30136586E-02	5.69169716	4.70311095 E-02	-6.59092116 E-04	
3(cubic	5.98388456E-03	5.62617288	6.50135886E-02	-1.94517664E-03	2.58649476E-05

Table 3: Section 5.1. Fitted polynomials, their coefficients and the corresponding RMSD using synthetic data. Observations up to 30 days.

Station	Polynomial	RM	ISD	<i>C</i> -	<i>C</i> .	<u>6-</u>	C-
Station	1 orynolliai	training	testing		c1	02	<i>C</i> 3
В	linear	1.16720184E-02	1.94601125 E-02	7.35751723	3.46560681E-02		
20	quadratic	2.20001211E-03	5.41537380E-03	7.30850076	4.32306214E-02	- 2.86981705E-04	
	cubic	1.67445669E-03	3.33740105E-03	7.29464777	4.71442613E-02	- 5.87655387E-04	6.73839556E-06
В	linear	3.07334766E-02	4.72819963E-02	5.81094708	2.65408393E-02		
000	quadratic	7.34536455E-03	2.00261102 E-02	5.68333537	4.88642187 E-02	- 7.47141135E-04	
30	cubic	3.07346962E-03	1.08396889E-02	5.61856923	6.71614492 E-02	- 2.15286467E-03	3.15036593E-05

Table 4: Section 5.1. Fitting polynomials, their coefficients, and the corresponding RMSD using synthetic data. In this case, observations of the first 10 days were used as training data to fit the polynomials. Observations of the remaining 20 days were considered unknown in the fitting phase and then they were used to test predictions given by the fitted polynomials.

Experiment 1: We assume that observed elevations correspond to instants $t_1 = 9.5$ and $t_2 = 10$ days



Figure 6: Section 5.1. Synthetic observed elevations at a given station and their approximations as (linear, quadratic, and cubic) polynomials of the inlet discharge. Observations up to 30 days were used to fit the polynomials

and 26 equally spaced stations between $x_{\min} = 0$ and $x_{\max} = 3000$ meters. We consider that the inlet discharge $Q_{\min}(t)$ at times t_1 and t_2 are also observed. We employ the model (5) with



Figure 7: Section 5.1. Synthetic observed elevations at a given station and their approximation as a (linear, quadratic, and cubic) polynomial of the inlet discharge. In this case, observations of the first 10 days were used as training data to fit the polynomials. Observations of the remaining 20 days were considered unknown in the fitting phase and, then, they were used to test predictions produced by the fitted polynomials.

q = 1 and s = p - 1 = 25. Note that, due to Theorem 2.1, it is not necessary to fit explicitly a polynomial with degree 25 in order to obtain predictions for the future at the given stations. Using this fitting, and considering suitable forecasts for the inlet discharges, we can predict elevations for days $11, 12, 13, \ldots, 29$ for 101 values of x equally spaced between x_{\min} and x_{\max} and we can compare these predictions with the observed elevations. Note that, in this case, the RMSD-error corresponding to the training set is necessarily equal to 0. The result of this experiment is given in Table 9.

- **Experiment 2:** Observed elevations correspond to instants $t_1 = 9, t_2 = 9.5$, and $t_3 = 10$ days. Elevation data correspond to these instances and the model (5) uses q = 2 and p - 1 = 25. So, the elevation at each station is modelled by a quadratic interpolating polynomial. The result of this experiment is given in Table 10.
- **Experiment 3:** Observed elevations correspond to instants $t_1 = 8.5, t_2 = 9, t_3 = 9.5$, and $t_4 = 10$ days. Elevation data correspond to these instances and the model (5) uses q = 3 and p 1 = 25. So, the elevation at each station is modelled by a cubic interpolating polynomial. The result of this experiment is given in Table 11.
- **Experiment 4:** Observed elevations correspond to instants $t \in \{3.5, 4, 4.5, \dots, 10\}$ days. Elevation data correspond to these instances and the model (5) is a line that fits the observed elevations at those instants in the least-squares sense. The result of this experiment is given in Table 12.
- **Experiment 5:** Observed elevations correspond to instants $t \in \{3.5, 4, 4.5, ..., 10\}$ days. Elevation data correspond to these instances and the model (5) is a quadratic polynomial that fits the observed elevations at those instants in the least-squares sense. The result of this experiment is given in Table 13.
- **Experiment 6:** Observed elevations correspond to instants $t \in \{3.5, 4, 4.5, ..., 10\}$ days. Elevation data correspond to these instances and the model (5) is a cubic polynomial that fits the observed elevations at those instants in the least-squares sense. The result of this experiment is given in Table 14.

Tables 9–14 in the Appendix show the results. Figures 8 and 9 give a graphical representation of the predictions' RMSD as a function of $t \in \{11, 12, ..., 29\}$. For each t, the RMSD of the 26 equally spaced $x \in [0, 3000]$ meters is shown. The experiment shows that polynomial interpolators of past data are bad at extrapolating to predict the future. One reason may be that they are based on little data and focus on capturing local behavior. Thus, the linear and quadratic options are less bad than the cubic, which quickly goes to infinity under the influence of local behavior. On the other hand, least squares polynomials computed with more data better capture the trend implicit in the data and thus better predict the future. Of the three options (linear, quadratic, and cubic), the cubic provides the best predictions.

5.3 Next-day predictions using observations on a mesh

In this experiment, we evaluate the six approaches considered in the previous subsection to predict the "elevation of the next day". We consider $t_{today} \in \{2, 3, ..., 28\}$ days and $t_{tomorrow} = t_{today} + 1$. Available data of z(x,t) with t multiple of half day and $t \leq t_{today}$ was used as training data. For the interpolating polynomials, only the most recent information was considered, while for least squares, all available data was considered. For each of the 26 equally spaced stations x between $x_{\min} = 0$ and $x_{\max} = 3000$ meters, the six approaches were used to predict the elevation $z(x, t_{tomorrow})$. Table 5 shows the details. As seen in previous experiments, least squares polynomials gave very reasonable



Figure 8: Section 5.2. RMSD of predictions of z(x,t) for $t \in \{10, 11, \ldots, 29\}$ when predictions are given by interpolating polynomials (linear, quadratic, and cubic) computed using training data with t < 10. For each t, the RMSD of the 26 equidistant $x \in [0, 3000]$ meters is being displayed.

predictions (with an average error of 1 centimeter in the case of the cubic polynomial) and performed better than interpolating polynomials. As expected, the cubic was better than the quadratic, which was better than the linear. Unlike previous experiments, interpolating polynomials were also useful in many cases, because in the present experiments we are dealing with next-day predictions, i.e. interpolating polynomials are used to extrapolate only a little outside the interpolating range.

5.4 Next-day predictions using irregularly distributed data

In the experiments of the previous subsection, we considered observations every 12 hours between day t = 3 and day t = 10 (15 time instants) at 26 stations equidistant between 0 and 3000 meters, totalizing 390 observations. However, considering our synthetic data, in that same domain of space (x, t) we have available data from hour to hour and at 101 equidistant stations, amounting to $101 \times 169 = 17069$ available data. With the intuition of using random subsets of data with uniform distribution, in the next experiment we draw the observations among the available data with probability $\frac{390/17069}{\nu} \approx \frac{0.0288}{\nu}$, with $\nu \in \{1, 2, 4\}$. With this way of determining the observations, we constituted training data sets with 394, 195, and 95 elevation observations.

The experiment consists of (a) positioning n_{stat} stations using Algorithm 4.1, (b) with the stations already positioned solve the linear least squares problem (18) that computes the cubic polynomial of each station, and (c) use those polynomials to predict the elevation of the "next day", that is, the day $t_{\text{tomorrow}} = 11$ at 101 equidistant points between 0 and 3000 meters. We wish to understand how the predictions behave for different values of n_{stat} .

Table 6 shows the results when 394 observations are available with the number of virtual stations varying from 2 to 100. The first column shows the number of stations. The second column reports



Figure 9: Section 5.2. RMSD of predictions of z(x,t) for $t \in \{10, 11, \ldots, 29\}$ when predictions are given by best fitting polynomials (linear, quadratic, and cubic) computed by solving a linear least squares problem using training data with $t \in \{3.5, 4, 4.5, \ldots, 9.5\}$. For each t, the RMSD of the 26 equidistant $x \in [0, 3000]$ meters is being displayed.

"minobs", the minimum number of observations that were used, given the positions of the virtual stations, to determine each of the n_{stat} cubic polynomials by means of least-square calculations. The third column shows the RMSD of the training data. The last column shows the RMSD of the next-day prediction at the 101 points equidistant between 0 and 3000 meters. It is clear from the figures in the table that the RMSD of the training data decreases monotonically as the number of stations increases. On the other hand, the RMSD of the next-day prediction remains more or less constant (between 3 and 6 centimeters) when the number of virtual stations is between 2 and 49 and deteriorates rapidly when this number is 50 or more. In fact, the optimal number of virtual stations is, in this case, 19. For completeness, in this case we report the results up to 100 virtual stations.

In Table 7 and Table 8 we report the same type of results when the number of available observations is 185 and 95, respectively. In the first case, the number of virtual stations goes from 2 to 37 and in the second case it goes from 2 to 19 because larger numbers of virtual stations yield prediction errors that are bigger than 1 meter. Again, the error in the training set decreases with the number of virtual stations, as the number of free parameters is increased.

6 Conclusions

This paper discusses the potential of methods based on surface elevation data alone for predicting river levels, provided that reliable inlet discharge forecasts $Q(x_{\min}, t)$ are available. We have focused on low-degree polynomial models because they are simple and economical in terms of the number of unknown parameters. The various alternatives presented in this paper can be considered successful in the sense that they provide results that are accurate enough for predicting the levels of real rivers.

	In	terpolati	ng		Fitting	
$t_{\rm today}$	polyno	omial of a	legree:	polyno	omial of o	legree:
	1	2	3	1	2	3
2	0.0062	0.0069	0.0199	0.0141	0.0021	0.0199
3	0.0130	0.0679	0.4752	0.0421	0.0219	0.0426
4	0.0051	0.0011	0.0033	0.0302	0.0108	0.0057
5	0.0014	0.0154	0.0348	0.0329	0.0142	0.0043
6	0.0086	0.0120	0.0150	0.0248	0.0094	0.0022
7	0.0071	0.0071	0.0071	0.0266	0.0105	0.0052
8	0.0021	0.0005	0.0141	0.0176	0.0028	0.0047
9	0.0035	0.0038	0.0041	0.0303	0.0075	0.0030
10	0.0163	0.0295	0.5743	0.0461	0.0225	0.0055
11	0.0083	0.0018	0.0056	0.0035	0.0024	0.0056
12	0.0008	0.0030	0.0034	0.0133	0.0013	0.0037
13	0.0905	0.0246	1.1340	0.0357	0.0131	0.0053
14	0.0199	0.0089	0.0180	0.0043	0.0113	0.0081
15	0.0162	0.0038	0.0069	0.0484	0.0079	0.0010
16	0.0063	0.0122	0.0163	0.0600	0.0165	0.0059
17	3.7608	1.0794	0.2602	0.0075	0.0045	0.0021
18	0.0290	0.0174	3.9772	0.0297	0.0090	0.0008
19	0.0168	0.0198	0.0503	0.0391	0.0057	0.0012
20	0.0104	0.0121	0.0124	0.0405	0.0046	0.0021
21	0.0281	0.0142	0.3077	0.0299	0.0129	0.0055
22	0.0353	0.0331	0.0751	0.0174	0.0052	0.0060
23	0.0173	0.0048	0.0022	0.0656	0.0194	0.0076
24	0.0018	0.0015	0.0014	0.0635	0.0199	0.0094
25	0.0126	0.7339	3.3581	0.0019	0.0079	0.0043
26	0.0806	0.1488	0.2467	0.0230	0.0113	0.0035
27	0.0224	0.0517	0.3592	0.0396	0.0099	0.0030
28	0.0075	0.0044	0.0029	0.0430	0.0107	0.0054
	0.7243	0.2537	1.0417	0.0353	0.0118	0.0102

Table 5: Section 5.3. For a given day t_{today} and a given experiment (interpolating or fitting polynomial of degree 1, 2, or 3) the table shows the RMSD of the next-day predicted elevation of all 26 stations. The last row shows the overall RMSD of each approach.

In particular, the strategy of virtual stations presented in this paper seems to be useful in the case where observations are irregularly distributed. Moreover, this strategy preserves the best third-order polynomial approximations in the regularly distributed case.

It is interesting to consider the problem of predicting flow-rates Q(x, t) from elevation observations z(x, t) only. From the mass-conservation equation we have that

$$\frac{\partial A}{\partial t} + \frac{\partial Q}{\partial x} = 0.$$

Therefore,

$$Q(x,t) = Q(x_{\min},t) - \int_{x_{\min}}^{x} \frac{\partial A}{\partial t}(\xi,t)d\xi.$$

Virtual stations	minobs	RMSD training	RMSD testing	Virtual stations	minobs	RMSD training	RMSD testing
2	394	5.3283820057185995E-002	6.1583164478249554E-002	52	10	8.0329354317997658E-003	0.12968356044883630
3	194	3.2648300820944318E-002	4.3726318071762964E-002	53	10	8.0306312766861461E-003	0.12972255677477529
4	101	3.2573813718738569E-002	4.5187290605368642E-002	54	10	8.0187525830073773E-003	0.12981398947867417
5	96	3.2349694130166182E-002	4.6027937601361361E-002	55	9	8.0179333149921674E-003	0.12982155115549487
6	96	3.1639597470529801E-002	4.8685500605615238E-002	56	9	8.0079086981791198E-003	0.12997349516966566
7	96	2.2853081231014000E-002	3.5410636986479435E-002	57	9	7.9116879215693353E-003	0.13080885606219347
8	47	2.2333106995790799E-002	3.4395169668978959E-002	58	9	7.1120982718531822E-003	0.12975686606923589
9	47	2.2005150036587609E-002	3.5766673494248759E-002	59	9	7.0527831979785103E-003	0.12960764964637364
10	47	1.9266463631366117E-002	3.4100377670761683E-002	60	9	6.8200629596160949E-003	0.12970197134584049
11	47	1.8330036730951786E-002	3.8099252360178991E-002	61	9	6.8099002730974117E-003	0.13443085010705041
12	47	1.8247387283753448E-002	3.9287689605826834E-002	62	9	6.7569652637862266E-003	0.13432463882639850
13	47	1.6544563006014871E-002	3.6201727385641792E-002	63	8	6.7551245197326791E-003	0.13450127671524303
14	46	1.6109150476189795E-002	3.4048905093485481E-002	64	8	6.7356313704111668E-003	0.14227538180090210
15	46	1.5513927451101283E-002	3.2651624306286674E-002	65	8	6.7297367192108533E-003	0.15228129547423630
16	26	1.5337223488082084E-002	3.2449894571859761E-002	66	8	6.7266867935987414E-003	0.71829333520509586
17	26	1.5158717471376091E-002	3.2345999462907726E-002	67	3	6.7264323627948203E-003	0.71833153306163988
18	13	1.5156034949535248E-002	3.2305120343670203E-002	68	3	4.7223533876023351E-003	0.71808112648009770
19	13	1.5034755623023519E-002	3.1751814157051458E-002	69	3	4.6359280777656803E-003	0.80170172105254001
20	13	1.5018746780348795E-002	3.2280097922278560E-002	70	3	4.6330920332480260E-003	0.84002240218473490
21	13	1.4923539265815706E-002	3.3525846097928733E-002	71	3	4.6122345451842994E-003	0.83973278380279726
22	13	1.4859547396657539E-002	3.4453557117947238E-002	72	3	4.6102751068605808E-003	1.5559016307267453
23	13	1.3691065318011980E-002	5.3843426618723260E-002	73	3	4.5875787632858409E-003	1.5560072582416900
24	13	1.3659881827357711E-002	5.4143100409803191E-002	74	3	4.4777490065165777E-003	1.5562018595829066
25	13	1.3263279056685832E-002	5.3164207679156618E-002	75	3	4.4755600778543063E-003	1.6023004097916536
26	13	1.3175509007241182E-002	5.2752211202618811E-002	76	3	4.3307476396778153E-003	1.6026236880641358
27	13	1.2893886036635917E-002	5.3294152115026458E-002	77	3	3.9620225228191950E-003	1.6027498818868846
28	13	1.2272901759381408E-002	5.3934282047854318E-002	78	3	2.8886317493116574E-003	1.6104032246579405
29	13	1.2090872110758055E-002	5.5372017107663166E-002	79	3	2.8869280012161587E-003	1.6211332816669319
30	13	$1.2043880504159150 {\rm E}{\rm -}002$	5.6348350230236106E-002	80	3	2.8599432980772987E-003	1.8984411394933036
31	13	1.1923895103591405E-002	5.6647214439744825 E-002	81	3	2.7391034801038162E-003	1.8988767221527962
32	13	1.1864076693219016E-002	5.6304646204128742E-002	82	3	2.5652606094846665E-003	1.9160614771120590
33	13	1.1682392020077523E-002	5.7208920168502021E-002	83	3	2.5267865337923233E-003	1.9168010751012954
34	13	1.1670097881090850E-002	$5.6984529825676034 \mathrm{E}{-}002$	84	3	2.3775396132823608E-003	2.3251907404973591
35	13	1.1628659265365073 E-002	$5.5947727821659403 {\rm E}{\rm -}002$	85	3	2.3001696336729283E-003	2.3255810361822524
36	13	1.0416666600669637E-002	3.7676481978826011E-002	86	3	2.2706294451196921E-003	2.3784129971256784
37	13	1.0399791281676873E-002	3.7889247768345305E-002	87	3	2.1418982774388468E-003	2.3898547222289541
38	13	1.0383501483021772E-002	3.7653477041362522E-002	88	3	2.1403856079314533E-003	2.3907039892197419
39	13	9.9227025361913641E-003	3.8432530450767118E-002	89	3	2.1403856079314563E-003	2.4025492211247181
40	13	9.8817703068327638E-003	3.7475166481987364E-002	90	3	2.1403856079314503E-003	2.4257624451134867
41	12	9.5545957721222922E-003	$3.8662642310044168 {\rm E}{\rm -}002$	91	3	$1.2241141853634287 {\rm E}{\rm -}003$	2.7471502163021850
42	12	9.5259047134240975 E-003	$3.8864212195983724 {\rm E}{\rm -}002$	92	3	1.1935575705138712E-003	2.7739273204290273
43	12	9.5068693663715384E-003	3.9334825174204911E-002	93	3	1.1935575705138105E-003	2.7914574146435789
44	12	9.4031730453848442E-003	$4.0138579708408451 {\rm E}{\rm -}002$	94	3	1.1935575705138755E-003	10.769455041417071
45	11	9.3878986230084352 E-003	$4.0381481232975115 {\rm E}{\rm -}002$	95	3	$1.1935575705138961 {\rm E}{\rm -}003$	13.194184220825461
46	11	$9.2827157053270003 {\rm E}{\rm -}003$	$4.1040039457329369 {\rm E}{\text{-}}002$	96	3	$1.1919087363459152 {\rm E}{\rm -}003$	15.209373170340974
47	11	$9.1301445568780712 {\rm E}{\rm -}003$	$4.3543218230213378\mathrm{E}{-}002$	97	3	$1.0625333177912046 {\rm E}{\rm -}003$	16.484479467068468
48	11	$8.4370965844998008 {\rm E}{\rm -}003$	$4.3316691550977185 {\rm E}{\rm -}002$	98	3	$1.0625333177911834 {\rm E}{\rm -}003$	18.255510661328792
49	11	$8.4166324951854554 {\rm E}{\text{-}}003$	$4.3061145849278733 {\rm E}{\rm -}002$	99	3	$1.0625333177911502 {\rm E}{\rm -}003$	18.414545717188812
50	11	$8.2245582976612011 {\rm E}{\text -}003$	0.12923190100341689	100	3	$1.0625333177912118 {\rm E}{\rm -}003$	18.510711323780274
51	10	8.2013285286650847E-003	0.12971323080775976				

Table 6: Section 5.4. 394 random observations in the first 10 days. Effect of increasing the number of virtual stations. Reporting training and test.

So, if we have a good approximation for A(x,t), then we can obtain, in principle, a good approximation of Q(x,t) [6]. Moreover, according to the results of the present paper, a good approximation of z(x,t)can be obtained using elevation observations and Q_{\min} forecasts. Unfortunately, the cross wetted area A(x,t) can be obtained from z(x,t) only if we already know the bed elevation $z_b(x)$ and the geometric characteristics of the channel. This is the information that we have considered uncertain, and whose use we have tried to avoid above under the "only elevation" approach of the present paper. Therefore, predicting flow rates from data alone is more problematic than predicting surface elevations, and this issue deserves future study.

Virtual stations	minobs	RMSD training	RMSD testing
2	185	5.1248811463696108E-002	7.4089510567510924E-002
3	91	3.2529197002708163E-002	3.5519815707586833E-002
4	47	3.1284750173733118E-002	4.3991827708239235E-002
5	46	2.9878458008594868E-002	8.3143696581750387E-002
6	46	2.2450370975549642 E-002	6.8068533438989373E-002
7	22	2.1764814785424559E-002	6.5415859677875707E-002
8	22	2.0766831805240742 E-002	6.7534463202462633E-002
9	22	2.0573231136337876E-002	6.6790025825457275E-002
10	22	2.0104363353167974 E-002	6.3158699486610168E-002
11	22	1.8826320026367610 E-002	7.3860255266057703E-002
12	22	1.8270195922934992 E-002	9.1669436715929684E-002
13	22	1.5006593710845378E-002	7.7134955384989309E-002
14	10	$1.4856901876581540 {\rm E}{\rm -}002$	7.7603153454404841E-002
15	10	1.4558586705511977 E-002	7.0383131196731591E-002
16	10	1.4367127212092557 E-002	7.2430701225312186E-002
17	10	1.4056078954309907 E-002	7.6150540705157130E-002
18	10	1.3944065662380992 E-002	7.3380393469394803E-002
19	10	1.2663706268095641 E-002	9.2371846936037255E-002
20	10	1.2527967989703639E-002	9.3799256133081668E-002
21	10	1.2452824633423691 E-002	9.2679586134847808E-002
22	10	1.2430540420894570 E-002	9.2397760084283950E-002
23	8	1.2308000711774588E-002	0.10149722107077552
24	8	1.2170734928201659 E-002	0.10745970660633239
25	8	1.0773903953212724 E-002	0.11569597875671561
26	8	1.0703768079820394 E-002	0.12342432166878507
27	8	8.0880895546961377E-003	0.13473581683984887
28	5	8.0856468697171491E-003	0.13471641725588770
29	5	8.0624368937640255E-003	0.13730815995101689
30	5	7.5576773867298457E-003	0.14055458227230450
31	5	7.4451793362722103E-003	0.14283338653559183
32	5	7.3306658105015731E-003	0.15803100201596343
33	5	6.9433436122286933E-003	0.20812431672514720
34	5	6.7419124067725706E-003	0.84362496442554258
35	5	6.7171350118954724 E-003	0.84673106438894241
36	5	5.1430668400310421E-003	0.89984696571610800
37	5	2.7526193304603067 E-003	1.5797211827586006

Table 7: Section 5.4. 185 random observations in the first 10 days. Effect of increasing the number of virtual stations. Reporting training and test RMSD.

References

 A. Agresta, M. Baioletti, C. Biscarini, F. Caraffini, A. Milani, and V. Santucci, Using optimisation meta-heuristics for the roughness estimation problem in river flow analysis, *Applied Sciences* 11, 10575, 2021.

Virtual stations	minobs	RMSD training	RMSD testing
2	95	5.1699965039000095E-002	6.3173540327799205E-002
3	47	3.1098552914075955E-002	3.7681240571308741E-002
4	23	3.0008346683687421 E-002	5.4680180663268317 E-002
5	23	2.7343428686985052E-002	0.10084121692806179
6	13	2.6263510371117418 E-002	0.12687890844372421
7	13	2.5469694685872728 E-002	0.14131992067387011
8	11	2.5009773875898412 E-002	0.14894525241456291
9	11	1.8148205210940294 E-002	9.3893818673443930E-002
10	11	1.7516836515265203E-002	0.13792608105500823
11	11	1.4868146929722386 E-002	0.14102668906719881
12	7	$1.4669480794577859 {\rm E}{\rm -}002$	0.14028858123690985
13	7	1.4344105609466579 E-002	0.15561454960887050
14	7	1.3939622987893766E-002	0.17476910376467428
15	5	1.3899191317637099 E-002	0.17136513895286537
16	5	1.3641822085803633E-002	0.17664288679815060
17	5	8.7311056846887617E-003	0.55468488663305626
18	5	8.4229878615030840E-003	0.55920223834725657
19	5	3.6141721204698040E-003	2.0837223835016192

Table 8: Section 5.4. 95 random observations in the first 10 days. Effect of increasing the number of virtual stations. Reporting training and test RMSD.

- [2] M. T. Ayvaz, A linked simulation-optimization model for simultaneously estimating the Manning's surface roughness values and their parameter structures in shallow water flows, *Journal of Hydrology* 500, pp. 183–199, 2013.
- [3] M. Kh. Askar and K. K. Al-jumaily, A nonlinear optimization model for estimating Manning's roughness coefficient, in *Proceedings of the Twelfth International Water Technology Conference*, *IWTC12*, Alexandria, Egypt, 2008, pp. 1299–1306.
- [4] E. G. Birgin, M. R. Correa, V. A. González-López, J. M. Martínez, and D. S. Rodrigues, Randomly supported models for the prediction of flows in channels, Technical Report of CRIAB-Mathematical Engineering, *submitted*.
- [5] E. G. Birgin and J. M. Martínez, Accelerated derivative-free nonlinear least-squares applied to the estimation of Manning coefficients, *Computational Optimization and Applications* pp. 689–715, 2022.
- [6] E. G. Birgin and J. M. Martínez, A PDE-informed optimization algorithm for river flow predictions, *Numerical Algorithms*, to appear (DOI: 10.1007/s11075-023-01647-1).
- [7] Y. Ding, Y. Jia, and S. S. Y. Wang, Identification of Manning's roughness coefficients in shallow water flows, *Journal of Hydraulic Engineering*, pp. 501–510, 2004.
- [8] Y. Ding and S. S. Y. Wang, Identification of Manning's roughness coefficients in channel network using adjoint analysis, *International Journal of Computational Fluid Dynamics* 19, pp. 3–13, 2005.

- [9] W. W. Emmett, W. W. Myrick, and R. H. Meade, Field data describing the movement and storage of sediment in the East Fork River, Wyoming, Part 1, *River Hydraulics and Sediment Transport*, Report No. 1, 1979.
- [10] G. H. Golub and Ch. F. Van Loan, *Matrix Computations*, The Johns Hopkins University Press, 2013.
- [11] K. Guta and K. S. H. Prasad, Estimation of open channel flow parameters by using optimization techniques, *International Journal of Science and Research* 6, pp. 1295–1304, 2018.
- [12] Y. Jia and S. S. y. Wang, 2001, CCHE2D: Two-dimensional Hydrodynamic and Sediment Transport Model for Unsteady Open Channel Flows Over Loose Bed, National Center for Computational Hydroscience and Engineering Technical Report No.: NCCHE-TR-2001-1, 2001.
- [13] R. J. LeVeque, Numerical Methods for Conservation Laws, Lectures in Mathematics, ETH Zürich, Birkäuser, 1992.
- [14] W. A. Marcus, K. Roberts, L. Harvey, and G. Tackman, An evaluation of methods for estimating Manning's n in small mountain streams, *Mountain Research and Development* 12, pp. 227–239, 1992.
- [15] J. M. Martínez and L. T. Santos, Inexact-restoration modelling with monotone interpolation and parameter estimation, *Optimization and Engineering*, to appear (DOI: 10.1007/s11081-023-09861-5).
- [16] R. H. Meade, W. W. Myrick, and W. W. Emmett, Field data describing the movement and storage of sediment in the East Fork River, Wyoming, Part 2, *River Hydraulics and Sediment Transport*, Report No. 2, 1979.
- [17] F. Pappenberger, K. Beven, M. Horrit, and S.Blazkova, Uncertainty in the calibration of effective roughness parameters in HEC-RAS using inundation and downstream level observations, *Journal* of Hydrology 302, pp. 46–69, 2005.
- [18] A. P. Piotrowski and J. J. Napiorkowski, Optimizing neural networks for river flow forecasting - Evolutionary Computation methods versus the Levenberg-Marquardt approach, *Journal of Hydrology* 407, pp. 12–27, 2011,
- [19] R. M. Porto, *Hidráulica Básica*, EESC-USP, São Paulo, 2000.
- [20] A. J. C. Saint-Venant, Théorie du mouvement non-permanent des eaux, avec application aux crues des rivière at à l'introduction des marées dans leur lit, *Comptes Rendus des Séances de* Académie des Sciences 73, pp. 147–154, 1871.
- [21] A. L. Simões, H. E. Schulz, and R. M. Porto, Métodos Computacionais em Hidráulica, Edufba, Bahia, 2017.

Appendix

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29	12.80	0.0648	0.0647	0.0615	0.0595	0.0712	0.0671	0.0644	0.0850	0.0930	0.0925	0.0988	0.0941	0.0973	0.0996	0.1009	0.1000	0.0901	0.0791	0.0963	0.1050	0.1095	0.1192	0.1167	0.1073	0.1053	0.0894	0.0932
28	11.80	0.0769	0.0765	0.0723	0.0691	0.0837	0.0786	0.0746	0.0997	0.1099	0.1091	0.1168	0.1109	0.1146	0.1174	0.1190	0.1178	0.1056	0.0919	0.1118	0.1219	0.1290	0.1414	0.1385	0.1276	0.1254	0.1073	0.1098
27	8.76	0.1213	0.1197	0.1102	0.1004	0.1261	0.1168	0.1059	0.1472	0.1661	0.1642	0.1764	0.1664	0.1702	0.1750	0.1771	0.1746	0.1536	0.1302	0.1559	0.1672	0.1894	0.2129	0.2085	0.1931	0.1906	0.1677	0.1634
26	6.54	0.1655	0.1627	0.1466	0.1278	0.1662	0.1527	0.1329	0.1884	0.2185	0.2150	0.2319	0.2178	0.2206	0.2283	0.2312	0.2271	0.1973	0.1643	0.1907	0.1973	0.2417	0.2793	0.2735	0.2544	0.2530	0.2272	0.2126
25	3.69	0.2494	0.2442	0.2142	0.1749	0.2407	0.2209	0.1819	0.2556	0.3113	0.3053	0.3317	0.3108	0.3089	0.3261	0.3327	0.3255	0.2795	0.2291	0.2492	0.2340	0.3279	0.4008	0.3926	0.3663	0.3718	0.3409	0.3018
24	3.76	0.2464	0.2412	0.2117	0.1730	0.2375	0.2179	0.1796	0.2526	0.3070	0.3010	0.3267	0.3060	0.3042	0.3205	0.3267	0.3196	0.2744	0.2248	0.2452	0.2313	0.3227	0.3930	0.3849	0.3591	0.3640	0.3337	0.2968
23	5.98	0.1785	0.1752	0.1568	0.1349	0.1770	0.1624	0.1397	0.1984	0.2317	0.2276	0.2456	0.2303	0.2325	0.2408	0.2438	0.2393	0.2072	0.1717	0.1973	0.2017	0.2525	0.2938	0.2876	0.2678	0.2668	0.2408	0.2241
22	10.60	0.0923	0.0916	0.0856	0.0803	0.0985	0.0919	0.0858	0.1165	0.1293	0.1281	0.1373	0.1300	0.1337	0.1371	0.1388	0.1372	0.1220	0.1051	0.1274	0.1384	0.1497	0.1655	0.1620	0.1496	0.1470	0.1273	0.1282
21	15.10	0.0438	0.0442	0.0432	0.0432	0.0509	0.0486	0.0482	0.0618	0.0672	0.0670	0.0717	0.0687	0.0718	0.0736	0.0749	0.0744	0.0683	0.0613	0.0743	0.0810	0.0821	0.0885	0.0869	0.0796	0.0783	0.0648	0.0689
20	13.40	0.0600	0.0601	0.0579	0.0568	0.0679	0.0643	0.0626	0.0819	0.0897	0.0894	0.0958	0.0914	0.0951	0.0975	0.0991	0.0983	0.0892	0.0790	0.0961	0.1048	0.1084	0.1179	0.1156	0.1062	0.1044	0.0879	0.0912
19	10.50	0.0948	0.0941	0.0882	0.0830	0.1021	0.0953	0.0891	0.1213	0.1350	0.1339	0.1438	0.1362	0.1404	0.1442	0.1461	0.1444	0.1286	0.1109	0.1343	0.1458	0.1583	0.1755	0.1719	0.1587	0.1563	0.1353	0.1348
18	6.62	0.1641	0.1613	0.1456	0.1272	0.1654	0.1521	0.1327	0.1881	0.2181	0.2147	0.2317	0.2176	0.2208	0.2286	0.2315	0.2275	0.1979	0.1649	0.1916	0.1984	0.2426	0.2804	0.2747	0.2555	0.2542	0.2281	0.2128
17	4.18	0.2316	0.2268	0.1999	0.1650	0.2246	0.2060	0.1712	0.2419	0.2920	0.2862	0.3105	0.2908	0.2903	0.3048	0.3101	0.3036	0.2610	0.2143	0.2361	0.2267	0.3099	0.3741	0.3665	0.3419	0.3455	0.3160	0.2824
16	4.98	0.2061	0.2020	0.1793	0.1509	0.2019	0.1851	0.1562	0.2219	0.2640	0.2589	0.2802	0.2625	0.2636	0.2749	0.2789	0.2734	0.2358	0.1942	0.2183	0.2162	0.2840	0.3369	0.3300	0.3077	0.3089	0.2810	0.2552
15	7.62	0.1419	0.1397	0.1271	0.1132	0.1444	0.1331	0.1182	0.1660	0.1895	0.1869	0.2010	0.1891	0.1923	0.1981	0.2004	0.1972	0.1722	0.1445	0.1707	0.1805	0.2119	0.2409	0.2358	0.2189	0.2165	0.1926	0.1848
14	12.70	0.0666	0.0665	0.0635	0.0615	0.0738	0.0696	0.0669	0.0884	0.0970	0.0965	0.1033	0.0983	0.1018	0.1044	0.1059	0.1049	0.0946	0.0831	0.1012	0.1104	0.1152	0.1256	0.1231	0.1132	0.1112	0.0944	0.0976
13	27.40	0.0030	0.0037	0.0046	0.0057	0.0063	0.0067	0.0077	0.0086	0.0093	0.0095	0.0104	0.0104	0.0116	0.0121	0.0125	0.0127	0.0128	0.0128	0.0145	0.0153	0.0144	0.0151	0.0151	0.0136	0.0138	0.0097	0.0113
12	25.50	0.0019	0.0021	0.0025	0.0030	0.0033	0.0034	0.0038	0.0044	0.0048	0.0048	0.0053	0.0052	0.0057	0.0060	0.0062	0.0062	0.0062	0.0061	0.0070	0.0074	0.0070	0.0074	0.0073	0.0066	0.0067	0.0049	0.0056
11	28.20	0.0041	0.0051	0.0065	0.0082	0.0090	0.0096	0.0110	0.0123	0.0133	0.0137	0.0149	0.0150	0.0167	0.0174	0.0181	0.0183	0.0186	0.0187	0.0210	0.0222	0.0209	0.0218	0.0218	0.0196	0.0200	0.0139	0.0163
1 days	$\min(t)$	0	120	240	360	480	600	720	840	960	1080	1200	1320	1440	1560	1680	1800	1920	2040	2160	2280	2400	2520	2640	2760	2880	3000	
t ii	0 ¹							_	_				ers	19u	ıu	į x		_			_			_			_	

Table 9: Section 5.2. For each value of x in the first column, predictions for t > 10 days use data observed at $t_1 = 9.5$ and $t_2 = 10$ days. Observed data correspond to $z(x, t_1)$ and $z(x, t_2)$ and the prediction is given by the *linear polynomial in* $Q_{\min}(t)$ that interpolates the data. Each cell of the table shows $|z_{\text{pred}}(x,t) - z(x,t)|$, where values of z(x,t) correspond to synthetic data that is not used in the prediction. The last line in the table shows the RMSD for each t.

29	12.80	0.0384	0.0462	0.0579	0.0731	0.0783	0.0834	0.0967	0.1059	0.1131	0.1157	0.1255	0.1261	0.1400	0.1455	0.1508	0.1528	0.1543	0.1548	0.1772	0.1884	0.1734	0.1792	0.1787	0.1606	0.1623	0.1139	0.1363
28	11.80	0.0437	0.0530	0.0671	0.0857	0.0909	0.0972	0.1135	0.1232	0.1308	0.1341	0.1452	0.1461	0.1625	0.1688	0.1750	0.1775	0.1799	0.1812	0.2075	0.2208	0.2014	0.2071	0.2066	0.1853	0.1872	0.1302	0.1583
27	8.76	0.0602	0.0753	0.0997	0.1327	0.1368	0.1478	0.1773	0.1884	0.1963	0.2019	0.2181	0.2206	0.2470	0.2559	0.2654	0.2699	0.2762	0.2809	0.3249	0.3488	0.3080	0.3118	0.3110	0.2780	0.2800	0.1898	0.2410
26	6.54	0.0683	0.0885	0.1239	0.1725	0.1726	0.1882	0.2320	0.2439	0.2483	0.2566	0.2764	0.2808	0.3168	0.3268	0.3390	0.3455	0.3564	0.3655	0.4288	0.4675	0.3991	0.3968	0.3958	0.3525	0.3532	0.2334	0.3099
25	3.69	0.0613	0.0896	0.1451	0.2241	0.2095	0.2321	0.3031	0.3190	0.3092	0.3214	0.3438	0.3518	0.4054	0.4116	0.4249	0.4355	0.4564	0.4749	0.5741	0.6495	0.5237	0.4976	0.4968	0.4403	0.4338	0.2712	0.3964
24	3.76	0.0622	0.0904	0.1453	0.2234	0.2097	0.2322	0.3022	0.3182	0.3094	0.3217	0.3444	0.3522	0.4054	0.4124	0.4260	0.4364	0.4567	0.4746	0.5727	0.6464	0.5234	0.4995	0.4987	0.4422	0.4364	0.2744	0.3967
23	5.98	0.0696	0.0913	0.1301	0.1836	0.1824	0.1993	0.2475	0.2603	0.2636	0.2727	0.2937	0.2987	0.3378	0.3481	0.3611	0.3682	0.3803	0.3903	0.4599	0.5037	0.4274	0.4235	0.4225	0.3762	0.3764	0.2479	0.3307
22	10.60	0.0508	0.0622	0.0799	0.1035	0.1088	0.1167	0.1376	0.1482	0.1565	0.1605	0.1738	0.1752	0.1953	0.2027	0.2102	0.2133	0.2169	0.2192	0.2518	0.2685	0.2425	0.2483	0.2476	0.2219	0.2240	0.1546	0.1903
21	15.10	0.0246	0.0293	0.0360	0.0447	0.0483	0.0512	0.0587	0.0648	0.0695	0.0711	0.0771	0.0773	0.0855	0.0889	0.0921	0.0932	0.0939	0.0938	0.1071	0.1137	0.1055	0.1094	0.1091	0.0981	0.0992	0.0700	0.0832
20	13.40	0.0335	0.0403	0.0502	0.0632	0.0675	0.0719	0.0833	0.0909	0.0969	0.0992	0.1074	0.1079	0.1198	0.1244	0.1289	0.1306	0.1322	0.1328	0.1516	0.1609	0.1478	0.1524	0.1520	0.1364	0.1379	0.0963	0.1165
19	10.50	0.0503	0.0618	0.0796	0.1033	0.1081	0.1162	0.1373	0.1470	0.1547	0.1587	0.1716	0.1732	0.1931	0.2003	0.2076	0.2109	0.2150	0.2178	0.2500	0.2667	0.2393	0.2440	0.2434	0.2179	0.2199	0.1505	0.1881
18	6.62	0.0677	0.0877	0.1226	0.1705	0.1704	0.1859	0.2291	0.2406	0.2448	0.2529	0.2722	0.2767	0.3121	0.3218	0.3337	0.3402	0.3511	0.3603	0.4226	0.4607	0.3927	0.3899	0.3888	0.3462	0.3468	0.2286	0.3051
17	4.18	0.0651	0.0920	0.1433	0.2160	0.2052	0.2266	0.2920	0.3067	0.3005	0.3123	0.3345	0.3419	0.3918	0.3997	0.4134	0.4231	0.4417	0.4580	0.5500	0.6169	0.5034	0.4838	0.4828	0.4283	0.4238	0.2685	0.3835
16	4.98	0.0685	0.0930	0.1383	0.2017	0.1958	0.2152	0.2723	0.2857	0.2843	0.2949	0.3167	0.3230	0.3675	0.3770	0.3905	0.3990	0.4145	0.4278	0.5091	0.5645	0.4685	0.4569	0.4558	0.4050	0.4030	0.2599	0.3599
15	7.62	0.0656	0.0832	0.1130	0.1534	0.1562	0.1695	0.2057	0.2177	0.2248	0.2317	0.2501	0.2535	0.2847	0.2946	0.3056	0.3111	0.3193	0.3257	0.3791	0.4096	0.3569	0.3591	0.3582	0.3198	0.3215	0.2162	0.2783
14	12.70	0.0383	0.0461	0.0578	0.0732	0.0781	0.0833	0.0968	0.1055	0.1124	0.1151	0.1248	0.1254	0.1393	0.1447	0.1499	0.1520	0.1538	0.1545	0.1767	0.1879	0.1723	0.1776	0.1771	0.1591	0.1607	0.1123	0.1355
13	27.40	0.0062	0.0071	0.0083	0.0098	0.0109	0.0114	0.0126	0.0145	0.0157	0.0160	0.0173	0.0172	0.0189	0.0197	0.0203	0.0205	0.0204	0.0200	0.0229	0.0244	0.0232	0.0243	0.0242	0.0219	0.0221	0.0160	0.0184
12	25.50	0.0019	0.0021	0.0025	0.0030	0.0033	0.0034	0.0038	0.0044	0.0048	0.0048	0.0053	0.0052	0.0057	0.0060	0.0062	0.0062	0.0062	0.0061	0.0070	0.0074	0.0070	0.0074	0.0073	0.0066	0.0067	0.0049	0.0056
11	28.20	0.0101	0.0116	0.0134	0.0159	0.0176	0.0183	0.0204	0.0233	0.0253	0.0257	0.0279	0.0278	0.0305	0.0317	0.0327	0.0330	0.0328	0.0322	0.0369	0.0393	0.0373	0.0391	0.0389	0.0351	0.0355	0.0257	0.0295
in days	$2_{\min(t)}$	0	120	240	360	480	009	720	840	960	1080	1200	1320	1440	1560	3 1680	1800	1920	2040	2160	2280	2400	2520	2640	2760	2880	3000	
t	اٽ ا												540	tən	ιu	. <i>x</i>												

Table 10: Section 5.2. For each value of x in the first column, predictions for t > 10 days use data observed at $t_1 = 9$, $t_2 = 9.5$, and $t_3 = 10$ days. Observed data correspond to $z(x, t_1)$, $z(x, t_2)$, and $z(x, t_3)$ and the prediction is given by the quadratic polynomial in $Q_{\min}(t)$ that interpolates the data. Each cell of the table shows $|z_{\text{pred}}(x,t) - z(x,t)|$, where values of z(x,t) correspond to synthetic data that is not used in the prediction. The last line in the table shows the RMSD for each t.

_	_	_																										
29	12.80	4.9138	5.6616	6.6283	7.8936	8.7966	9.1615	10.2165	11.7513	12.7579	12.9940	14.1174	14.0431	15.4328	16.0500	16.5833	16.7394	16.6051	16.3045	18.7609	19.9980	19.0078	19.9587	19.8726	17.9594	18.1309	13.2084	14.9921
28	11.80	6.3488	7.3148	8.5636	10.1980	11.3650	11.8363	13.1990	15.1824	16.4833	16.7884	18.2399	18.1439	19.9393	20.7368	21.4259	21.6275	21.4537	21.0649	24.2383	25.8364	24.5584	25.7875	25.6763	23.2045	23.4262	17.0667	19.3700
27	8.76	12.3497	14.2281	16.6552	19.8311	22.1029	23.0189	25.6665	29.5247	32.0571	32.6500	35.4732	35.2860	38.7771	40.3286	41.6687	42.0604	41.7215	40.9645	47.1322	50.2365	47.7585	50.1520	49.9359	45.1292	45.5611	33.1948	37.6696
26	6.54	18.5422	21.3614	25.0027	29.7662	33.1791	34.5538	38.5247	44.3152	48.1204	49.0096	53.2481	52.9667	58.2056	60.5356	62.5474	63.1350	62.6253	61.4882	70.7401	75.3927	71.6838	75.2812	74.9569	67.7422	68.3923	49.8317	56.5431
25	3.69	29.1420	33.5694	39.2832	46.7554	52.1257	54.2847	60.5135	69.6028	75.5906	76.9867	83.6463	83.2037	91.4274	95.0931	98.2557	99.1776	98.3739	96.5853	111.1036	118.3897	112.5881	118.2556	117.7459	106.4133	107.4409	78.2905	88.8180
24	3.76	28.8421	33.2241	38.8794	46.2752	51.5897	53.7266	59.8918	68.8878	74.8134	76.1951	82.7860	82.3479	90.4873	94.1148	97.2445	98.1570	97.3620	95.5921	109.9618	117.1742	111.4305	117.0381	116.5336	105.3176	106.3341	77.4831	87.9043
23	5.98	20.3765	23.4744	27.4749	32.7081	36.4590	37.9695	42.3320	48.6940	52.8761	53.8529	58.5103	58.2010	63.9570	66.5176	68.7282	69.3738	68.8137	67.5644	77.7289	82.8392	78.7655	82.7191	82.3627	74.4352	75.1499	54.7554	62.1302
22	10.60	8.4059	9.6847	11.3377	13.5008	15.0463	15.6701	17.4736	20.0996	21.8224	22.2262	24.1479	24.0206	26.3974	27.4534	28.3656	28.6324	28.4021	27.8873	32.0876	34.2026	32.5122	34.1400	33.9928	30.7205	31.0141	22.5951	25.6436
21	15.10	2.4647	2.8398	3.3249	3.9599	4.4128	4.5959	5.1254	5.8954	6.4002	6.5188	7.0823	7.0451	7.7424	8.0520	8.3196	8.3979	8.3305	8.1797	9.4124	10.0333	9.5362	10.0132	9.9700	9.0102	9.0962	6.6266	7.5214
20	13.40	4.1676	4.8018	5.6218	6.6951	7.4612	7.7707	8.6656	9.9678	10.8218	11.0221	11.9751	11.9121	13.0910	13.6146	14.0670	14.1994	14.0853	13.8300	15.9140	16.9637	16.1241	16.9311	16.8581	15.2352	15.3806	11.2054	12.7173
19	10.50	8.5959	9.9036	11.5939	13.8058	15.3867	16.0245	17.8686	20.5547	22.3169	22.7298	24.6952	24.5650	26.9957	28.0757	29.0087	29.2815	29.0456	28.5187	32.8142	34.9771	33.2496	34.9151	34.7646	31.4182	31.7185	23.1091	26.2250
18	6.62	18.2898	21.0707	24.6625	29.3613	32.7280	34.0839	38.0008	43.7130	47.4666	48.3438	52.5248	52.2471	57.4149	59.7135	61.6980	62.2776	61.7746	60.6527	69.7791	74.3685	70.7105	74.2596	73.9397	66.8231	67.4644	49.1561	55.7752
17	4.18	27.0883	31.2045	36.5174	43.4658	48.4563	50.4634	56.2557	64.7070	70.2716	71.5693	77.7600	77.3484	84.9949	88.4012	91.3405	92.1977	91.4514	89.7892	103.2885	110.0664	104.6687	109.9337	109.4599	98.9249	99.8788	72.7786	82.5683
16	4.98	23.9477	27.5875	32.2868	38.4332	42.8435	44.6183	49.7420	57.2164	62.1341	63.2816	68.7550	68.3912	75.1538	78.1641	80.7623	81.5206	80.8615	79.3924	91.3322	97.3308	92.5522	97.2032	96.7844	87.4691	88.3109	64.3475	73.0076
15	7.62	15.3205	17.6504	20.6602	24.5982	27.4170	28.5532	31.8360	36.6212	39.7637	40.4988	44.0008	43.7685	48.0982	50.0230	51.6852	52.1710	51.7504	50.8114	58.4593	62.3073	59.2371	62.2071	61.9390	55.9770	56.5133	41.1748	46.7243
14	12.70	5.0472	5.8152	6.8082	8.1078	9.0355	9.4102	10.4938	12.0706	13.1047	13.3473	14.5013	14.4250	15.8525	16.4865	17.0343	17.1946	17.0565	16.7475	19.2708	20.5416	19.5250	20.5020	20.4136	18.4484	18.6246	13.5685	15.3999
13	27.40	0.0884	0.1018	0.1192	0.1420	0.1582	0.1648	0.1838	0.2112	0.2293	0.2335	0.2537	0.2524	0.2773	0.2884	0.2980	0.3008	0.2985	0.2931	0.3372	0.3594	0.3415	0.3585	0.3569	0.3225	0.3256	0.2371	0.2694
12	25.50	0.0019	0.0021	0.0025	0.0030	0.0033	0.0034	0.0038	0.0044	0.0048	0.0048	0.0053	0.0052	0.0057	0.0060	0.0062	0.0062	0.0062	0.0061	0.0070	0.0074	0.0070	0.0074	0.0073	0.0066	0.0067	0.0049	0.0056
11	28.20	0.1884	0.2171	0.2542	0.3027	0.3372	0.3512	0.3917	0.4503	0.4888	0.4978	0.5408	0.5380	0.5912	0.6148	0.6353	0.6413	0.6363	0.6249	0.7188	0.7662	0.7280	0.7642	0.7609	0.6876	0.6942	0.5055	0.5743
t in days	$Q_{\min}(t)$	0	120	240	360	480	009	720	840	960	1080	1200	ers 1320	1440 1440	n 1560	1680 i	1800	1920	2040	2160	2280	2400	2520	2640	2760	2880	3000	

Table 11: Section 5.2. For each value of x in the first column, predictions for t > 10 days use data observed at $t_1 = 8.5$, $t_2 = 9$, $t_3 = 9.5$, and $t_4 = 10$ days. Observed data correspond to $z(x, t_1)$, $z(x, t_2)$, $z(x, t_3)$, and $z(x, t_4)$ and the prediction is given by the *cubic polynomial in* $Q_{\min}(t)$ that interpolates the data. Each cell of the table shows $|z_{\text{pred}}(x, t) - z(x, t)|$, where values of z(x, t) correspond to synthetic data that is not used in the prediction. The last line in the table shows the RMSD for each t.

		29	$\begin{array}{c} 0.0058\\ 0.0055\\ 0.0076\\ 0.0076\\ 0.0076\\ 0.0067\\ 0.0067\\ 0.0088\\ 0.0110\\ 0.0110\\ 0.0110\\ 0.0113\\ 0.0122\\ 0.0122\\ 0.0122\\ 0.0122\\ 0.0122\\ 0.0128\\ 0.0128\\ 0.0128\\ 0.0114\\ 0.0114\\ 0.0114\\ 0.0114\\ 0.0114\\ 0.0114\\ 0.0114\\ 0.0114\\ 0.0114\\ 0.0114\\ 0.0116\\ 0.0016\\ 0.0000\\$
		28	$\begin{array}{c} 0.0002\\ 0.0007\\ 0.0008\\ 0.0015\\ 0.0015\\ 0.0016\\ 0.0016\\ 0.0034\\ 0.0034\\ 0.0034\\ 0.0040\\ 0.0040\\ 0.0040\\ 0.0051\\ 0.0050\\ 0.0050\\ 0.0051\\ 0.0050\\ 0.0050\\ 0.0050\\ 0.0050\\ 0.0077\\ 0.0077\\ 0.0077\\ 0.0077\\ 0.0072\\$
10.0	$\begin{array}{c} 0.0076\\ 0.0057\\ 0.0059\\ 0.0059\\ 0.0065\\ 0.0062\\ 0.0072\\ 0.0072\\ 0.0072\\ 0.0066\\ 0.0066\\ 0.0066\\ 0.0065\\ 0.0065\\ 0.0065\\ 0.0065\\ 0.0065\\ 0.0065\\ 0.0065\\ 0.0065\\ 0.0072\\ 0.0073\\$	27	$\begin{array}{c} 0.0246\\ 0.0233\\ 0.0133\\ 0.0197\\ 0.0197\\ 0.0197\\ 0.0167\\ 0.0162\\ 0.0226\\ 0.0226\\ 0.0226\\ 0.0226\\ 0.0226\\ 0.0226\\ 0.0113\\ 0.0105\\ 0.0105\\ 0.0105\\ 0.0105\\ 0.0105\\ 0.0253\\ 0.0105\\ 0.0253\\ 0.0105\\ 0.0253\\$
9.5	$\begin{array}{c} 0.0115\\ 0.0103\\ 0.0003\\ 0.0003\\ 0.0011\\ 0.0069\\ 0.0112\\ 0.0126\\ 0.0126\\ 0.0126\\ 0.0126\\ 0.0126\\ 0.0125\\ 0.0005\\$	26	$\begin{array}{c} 0.0545\\ 0.0526\\ 0.0420\\ 0.0431\\ 0.0281\\ 0.0434\\ 0.03569\\ 0.0544\\ 0.0569\\ 0.0569\\ 0.0561\\ 0.0541\\ 0.0541\\ 0.0541\\ 0.0541\\ 0.0541\\ 0.0541\\ 0.0567\\ 0.0567\\ 0.0185\\ 0.0057\\ 0.00649\\ 0.00628\\ 0.00628\\ 0.00639\\ 0.00649$
0.0	$\begin{array}{c} 0.0036\\ 0.0036\\ 0.0035\\ 0.0031\\ 0.0031\\ 0.0038\\ 0.0050\\ 0.0053\\ 0.0053\\ 0.0053\\ 0.0053\\ 0.0053\\ 0.0056\\ 0.0056\\ 0.0058\\$	25	$\begin{array}{c} 0.1.200\\ 0.1.153\\ 0.0.592\\ 0.0.592\\ 0.0.586\\ 0.0.564\\ 0.0.886\\ 0.0.886\\ 0.0.886\\ 0.0.886\\ 0.0.581\\ 0.1.121\\ 0.1.121\\ 0.1.123\\ 0.1.123\\ 0.1.123\\ 0.1.123\\ 0.1.238\\ 0.1.2$
8.5	$\begin{array}{c} 0.0077\\ 0.0074\\ 0.0064\\ 0.0070\\ 0.0070\\ 0.0079\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0002\\$	24	$\begin{array}{c} 0.1175\\ 0.1177\\ 0.0577\\ 0.0577\\ 0.0545\\ 0.08545\\ 0.08541\\ 0.01197\\ 0.1197\\ 0.1197\\ 0.1163\\ 0.1163\\ 0.1163\\ 0.1163\\ 0.1163\\ 0.1163\\ 0.1163\\ 0.1172\\ 0.1189\\ 0.1172\\ 0.1189\\ 0.1172\\ 0.1261\\ 0.1172\\ 0.0250\\ 0.1172\\ 0.0250\\ 0.1172\\ 0.0250\\ 0.1172\\ 0.01369\\ 0.1476\\ 0.1450\\ 0.1456\\ 0.1$
8.0	$\begin{array}{c} 0.0084\\ 0.0082\\ 0.0075\\ 0.0076\\ 0.0078\\ 0.0071\\ 0.0113\\ 0.0112\\ 0.012\\ 0.$	23	$\begin{array}{c} 0.0639\\ 0.0639\\ 0.0648\\ 0.0313\\ 0.0418\\ 0.04513\\ 0.04513\\ 0.0448\\ 0.0616\\ 0.0619\\ 0.0616\\ 0.0611\\ 0.0613\\ 0.0613\\ 0.0576\\ 0.0613\\ 0.0576\\ 0.0513\\ 0.0728\\ 0.0728\\ 0.0728\\ 0.0728\\ 0.0778\\ 0.078$
ays 7.5	$\begin{array}{c} 0.0108\\ 0.0105\\ 0.0079\\ 0.0079\\ 0.00104\\ 0.00134\\ 0.00136\\ 0.01136\\ 0.01136\\ 0.01136\\ 0.01136\\ 0.01135\\ 0.01135\\ 0.01136\\ 0.01110\\ 0.01115\\ 0.01115\\ 0.01115\\ 0.01115\\ 0.01110\\ 0.00110\\ 0.0010\\ 0.0010\\ 0.0001\\ 0.0000\\ 0.000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.000\\ 0.000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.00$	tys 22	$\begin{array}{c} 0.0076\\ 0.0076\\ 0.0054\\ 0.0056\\ 0.0056\\ 0.0036\\ 0.0036\\ 0.0047\\ 0.0047\\ 0.0047\\ 0.0028\\ 0.0028\\ 0.0028\\ 0.0028\\ 0.0028\\ 0.0028\\ 0.0028\\ 0.0028\\ 0.0028\\ 0.0028\\ 0.0028\\ 0.00016\\ 0.00016\\ 0.0017\\ 0.0016\\ 0.0017\\ 0.0017\\ 0.0017\\ 0.0017\\ 0.0017\\ 0.0017\\ 0.0017\\ 0.0017\\ 0.0001\\ 0.00028\\ 0.00028\\ 0.00028\\ 0.00028\\ 0.00016\\ 0.00000\\ 0.00000\\ 0.000\\ 0.000\\ $
s - t in d 7.0	$\begin{array}{c} 0.0125\\ 0.0120\\ 0.01166\\ 0.01166\\ 0.01166\\ 0.01152\\ 0.01152\\ 0.01148\\ 0.0148\\ 0.0148\\ 0.0148\\ 0.0148\\ 0.0148\\ 0.0148\\ 0.0115\\ 0.0112\\ 0.0115\\ 0.0015\\ 0.0005\\ 0$	-t in de 21	$\begin{array}{c} 0.0119\\ 0.0115\\ 0.01081\\ 0.01081\\ 0.0114\\ 0.0114\\ 0.01157\\ 0.0157\\ 0.0157\\ 0.0156\\ 0.0156\\ 0.0156\\ 0.0156\\ 0.0167\\ 0.0163\\ 0.0123\\ 0.0013\\ 0.00$
servations 6.5	$\begin{array}{c} 0.0063\\ 0.0058\\ 0.0046\\ 0.0033\\ 0.0046\\ 0.0035\\ 0.0054\\ 0.0054\\ 0.0054\\ 0.0054\\ 0.0043\\ 0.0043\\ 0.0043\\ 0.0043\\ 0.0043\\ 0.0043\\ 0.0043\\ 0.0043\\ 0.0043\\ 0.0043\\ 0.0046\\$	edictions 20	$\begin{array}{c} 0.0068\\ 0.0065\\ 0.0062\\ 0.0062\\ 0.0062\\ 0.0062\\ 0.0085\\ 0.0085\\ 0.0085\\ 0.0085\\ 0.0087\\ 0.0087\\ 0.0087\\ 0.0087\\ 0.0087\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.00112\\ 0.0012\\ 0.0012\\ 0.0012\\ 0.0002\\ 0$
SD of ob 6.0	$\begin{array}{c} 0.0127\\ 0.0121\\ 0.0121\\ 0.0112\\ 0.00112\\ 0.00112\\ 0.00112\\ 0.00112\\ 0.01124\\ 0.0147\\ 0.0143\\ 0.0143\\ 0.0143\\ 0.0143\\ 0.0143\\ 0.0143\\ 0.0143\\ 0.0143\\ 0.0143\\ 0.0143\\ 0.0143\\ 0.0103\\ 0.0103\\ 0.0103\\ 0.0165\\ 0.01063\\ 0.0103\\ 0.0165\\ 0.01063\\ 0.00163\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.000\\ 0.000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.000\\ 0.000\\ 0.000\\ 0.0000\\ 0.0000\\ 0.00$	ISD of pr 19	$\begin{array}{c} 0.0093\\ 0.0074\\ 0.0078\\ 0.0078\\ 0.0078\\ 0.0076\\ 0.0076\\ 0.0098\\ 0.0098\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0090\\ 0.0093\\ 0.0003\\$
$\frac{RM}{5.5}$	$\begin{array}{c} 0.0117\\ 0.0117\\ 0.0099\\ 0.0095\\ 0.00054\\ 0.0074\\ 0.00117\\ 0.0113\\ 0.0113\\ 0.0113\\ 0.0112\\ 0.0103\\ 0.0103\\ 0.0103\\ 0.0103\\ 0.0103\\ 0.0103\\ 0.0103\\ 0.0103\\ 0.0117\\ 0.0127\\ 0.01127\\ 0.01127\\ 0.01127\\ 0.01111\\ 0.01112\\ 0.01111\\ 0.0111$	RN. 18	$\begin{array}{c} 0.0536\\ 0.0512\\ 0.0411\\ 0.0411\\ 0.0389\\ 0.0441\\ 0.0382\\ 0.0572\\ 0.0572\\ 0.0572\\ 0.0572\\ 0.0572\\ 0.0554\\ 0.0551\\ 0.0551\\ 0.0551\\ 0.0551\\ 0.0551\\ 0.0551\\ 0.0551\\ 0.0204\\ 0.0204\\ 0.0203\\ 0.0645\\ 0.0645\\ 0.066\\ 0.066\\ 0.06\\ 0.$
5.0	$\begin{array}{c} 0.0092\\ 0.0071\\ 0.0071\\ 0.0075\\ 0.0075\\ 0.0042\\ 0.0094\\ 0.0098\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.0092\\ 0.00112\\ 0.0044\\ 0.0091\\ 0.0091\\ 0.0044\\ 0.0092\\ 0.00112\\ 0.00112\\ 0.00107\\ 0.0007\\ 0.0$	17	0.1053 0.1010 0.0811 0.0520 0.0541 0.0763 0.0769 0.0769 0.0769 0.0769 0.0769 0.0769 0.0763 0.0793 0.0745 0.07
4.5	$\begin{array}{c} 0.0048\\ 0.0046\\ 0.0037\\ 0.0024\\ 0.0024\\ 0.0021\\ 0.0055\\ 0.0055\\ 0.0053\\ 0.0053\\ 0.0055\\ 0.0056\\ 0.0056\\ 0.0056\\ 0.0056\\ 0.0056\\ 0.0056\\ 0.0056\\ 0.0056\\ 0.0056\\ 0.0056\\ 0.0073\\$	16	$\begin{array}{c} 0.0850\\ 0.0814\\ 0.0814\\ 0.04253\\ 0.04053\\ 0.06936\\ 0.0879\\ 0.0879\\ 0.0879\\ 0.0879\\ 0.0879\\ 0.0879\\ 0.0879\\ 0.0879\\ 0.0873\\ 0.0873\\ 0.0873\\ 0.0873\\ 0.0873\\ 0.0853\\ 0.0853\\ 0.0924\\ 0.0749\\ 0.0749\\ 0.0749\\ 0.0749\\ 0.0749\\ 0.0749\\ 0.0749\\ 0.0749\\ 0.0749\\ 0.07103\\ 0.0720\\ 0.0720\\ 0.0088\\ 0.0088\\ 0.0749\\ 0.0749\\ 0.0708\\ 0.0088\\ 0.0$
4.0	0.0019 0.0018 0.0012 0.0012 0.0012 0.0012 0.0017 0.0015 0.0015 0.0015 0.0015 0.0013 0.	15	$\begin{array}{c} 0.0379\\ 0.0360\\ 0.01360\\ 0.01361\\ 0.01361\\ 0.02371\\ 0.02371\\ 0.02372\\ 0.03373\\ 0.0378\\ 0.03323\\ 0.03323\\ 0.03323\\ 0.03323\\ 0.03323\\ 0.03323\\ 0.03323\\ 0.03323\\ 0.03323\\ 0.03332\\ 0.03332\\ 0.0157\\ 0.03333\\ 0.01157\\ 0.03337\\ 0.0157\\ 0.03337\\ 0.0157\\ 0.03357\\ 0.03357\\ 0.0157\\ 0.03357\\ 0.0157\\ 0.03357\\ 0.0157\\ 0.03357\\ 0.01257\\ 0.03357\\ 0.01257\\ 0.03357\\ 0.01257\\ 0.03357\\ 0.00357\\ 0.00357\\ 0.00357\\ 0.00357\\ 0.00357\\ 0.00357\\ 0.00357\\ 0.00357\\ 0.00357\\ 0.00357\\ 0.00357\\ 0.00357\\ 0.00157\\ 0.00357\\ 0.00357\\ 0.00357\\ 0.00357\\ 0.00357\\ 0.00357\\ 0.00357\\ 0.00355\\ 0.00355\\ 0.00355\\ 0.00355\\ 0.00355\\ 0.00355\\ 0.00355\\ 0.00355\\ 0.00355\\ 0.00355\\ 0.00355\\ 0.00355\\ 0.00355\\ 0.00355\\ 0.00055\\ 0.0005\\ 0.0$
3.5	$\begin{array}{c} 0.0052\\ 0.0078\\ 0.0072\\ 0.0072\\ 0.0072\\ 0.0072\\ 0.0094\\ 0.0092\\ 0.0094\\ 0.0092\\ 0.0090\\ 0.0092\\ 0.0091\\ 0.0091\\ 0.0091\\ 0.0091\\ 0.0065\\ 0.0074\\ 0.0092\\ 0.0074\\ 0.0065\\ 0.0074\\ 0.0092\\ 0.0065\\ 0.0112\\ 0.0065\\ 0.0009\\$	14	$\begin{array}{c} 0.0046\\ 0.0046\\ 0.0041\\ 0.0034\\ 0.0037\\ 0.0038\\ 0.0079\\ 0.0079\\ 0.0078\\ 0.0086\\ 0.0083\\ 0.0086\\ 0.0083\\ 0.0086\\ 0.0092\\ 0.0092\\ 0.0093\\ 0.0093\\ 0.0093\\ 0.0093\\ 0.0093\\ 0.0093\\ 0.0093\\ 0.0093\\ 0.00126\\ 0.0074\\ 0.0072\\ 0.0012\\ 0.0072\\ 0.0002\\ 0.0012\\ 0.0002$
3.0	0.0253 0.02142 0.0144 0.0144 0.01144 0.01214 0.01235 0.0126 0.0270 0.0255 0.0255 0.0255 0.0255 0.0255 0.0255 0.0255 0.0255 0.0255 0.0255 0.0255 0.0255 0.0255 0.0255 0.0253 0.0253 0.0253 0.0253 0.0253 0.0253 0.0253 0.0253 0.03337 0.03337 0.03337 0.03337 0.0333 0.0333 0.0333	13	0.0207 0.0194 0.0125 0.01700 0.01750 0.01761 0.01770 0.01770 0.01770 0.01770 0.01770 0.01783 0.01170 0.01183 0.02183 0.02184 0.01917 0.01917 0.01917 0.01936 0.01937 0.01823 0.01136 0.01137 0.01136 0.01137 0.01137 0.01137 0.01137 0.01137 0.01137 0.01137 0.01137 0.01137 0.01137 0.02211 0.02214 0.0214
	0 120 240 360 480 600 600 600 1080 1080 1080 11200 11200 11200 11200 11200 11200 2160 22040 2200 200	12	$\begin{array}{c} 0.0036\\ 0.0088\\ 0.0068\\ 0.0068\\ 0.0068\\ 0.0068\\ 0.0068\\ 0.0068\\ 0.0063\\ 0.0072\\ 0.0072\\ 0.0072\\ 0.0072\\ 0.0059\\ 0.0072\\ 0.0073\\$
	statam ni x	11	$\begin{array}{c} 0.0248\\ 0.0231\\ 0.0145\\ 0.01145\\ 0.01199\\ 0.01097\\ 0.01037\\ 0.0212\\ 0.0225\\ 0.0225\\ 0.0223\\ 0.0213\\ 0.0213\\ 0.0213\\ 0.0213\\ 0.02148\\ 0.0126\\ 0.0212\\ 0.0212\\ 0.0212\\ 0.02248\\ 0.0125\\ 0.02248\\ 0.0125\\ 0.02248\\ 0.00224\\ 0.02248\\ 0.00224\\ 0.00$
			0 1240 240 240 660 600 600 1080 11200 11320 11320 1140 1140 11320 11320 11320 11320 11320 11320 11920 22640 22660 22640 22660 22640 22660 22660 22660 22660 22660 22660 22660 22660 226000 226000 22600000000
			x in meters

Table 12: Section 5.2. For each value of x in the first column, predictions for t > 10 days use data observed at $t \in \{3.5, 4, 4.5, \ldots, 10\}$ days. Observed data correspond to z(x, t) and the prediction is given by the *best fitting linear polynomial (solution of a linear least squares problem)*. Each cell of the table shows $|z_{\text{pred}}(x,t) - z(x,t)|$. In the left-hand part of the table, values of z(x,t) correspond to synthetic trainind data (used in the fitting), while in the right-hand part of the table z(x,t) correspond to synthetic (testing) data that is not being used in the fitting. The last line in the table shows the RMSD for each t.

																																29	0.0027	0.0039	0.0038	0.0036	0.0053	0.0050	0.0046	0.0069	0.0087	0.0087	0.0098	0.0094	1010 0	0.0108	0.0112	0.0111	0.0101	0.0088	0.0099	0.0101	0.0128	0.0152	0.0151	0.0141	0.0144	0.0124	0.0101
																																28	0.0000	0.0031	0.0029	0.0025	0.0039	0.0037	0.0032	0.0049	0.0065	0.0064	0.0073	0.0070	0.0074	0.0080	0.0083	0.0082	0.0074	0.0064	0.0069	0.0067	0.0094	0.0114	0.0114	0.0107	0.0110	0.0096	0.0075
10.0		0.0012	0.0014	0.0014	0.0017	0.0017	0.0017	0.0021	0.0025	0.00055	0200.0	0.0027	0.0026	0.0028	0.0030	0.0031	0.0021	100000	0.0029	0.0028	0.0030	0.0030	0.002.4	100000	0.0037	0.0037	0.0034	0.0035	0.0027	0.0027		27	0.001.0	2100.0	0.0008	0.0007	0.0001	0.0003	0.0016	0.0014	0.0007	0.0009	0.0011	0.0013	0.0019	0.0020	0.0022	0.0023	0.0027	0.0030	0.0048	0.0064	0.0038	0.0032	0.0033	0.0028	0.0026	0.0014	0.0026
9.5	2	0.0012	0.0008	0.0004	0.0009	0.0007	0.0002	0.0008	0.0013	0.0019	7100.0	0.0014	0.0012	0.0011	0.0012	0.0012	0.0019	710000	0.0009	0.0005	0.0002	0.0001	0.0011	1100.0	1.100.0	0.0017	0.0017	0.0018	0.0018	0.0012		26	0.0193	0.0117	0.0081	0.0020	0.0077	0.0065	0.0008	0.0030	0.0002	0.0082	0.0095	0.0083	0.0066	0.0081	0.0084	0.0077	0.0052	0.0026	0.0044	0.0128	0.0029	0.0100	0.0100	0.0100	0.0121	0.0137	0.0087
0.6	2	0.0006	0.000	0.0008	0.0004	0.0006	0.0014	0.0009	0.0005	0.0007	1000.0	0.0007	0.0008	0.0011	0.0011	0.0012	0.0013	01000	0100.0	0.0020	0.0027	0.0033	0.0018	010000	0.0012	0.0012	0.0010	0.0009	0.0000	0.0013		25	0.0462	0.0450	0.0335	0.0136	0.0351	0.0325	0.0147	0.0170	0.0403	0.0380	0.0442	0.0409	0.0344	0.0436	0.0473	0.0448	0.0360	0.0271	0.0057	0.0278	0.0236	0.0550	0.0543	0.0519	0.0619	0.0629	0.0411
s. r.:	5	0.0001	0.0003	0.0005	0.0005	0.0005	0.0007	0.0007	0.0007	0.0007	0.000	0.0007	0.0008	0.0009	0.0009	0.0009	0,000	0100.0	0100.0	0.0010	0.0013	0.0015	0.0019	7100.0	0.0011	0.0011	0.0010	0.0009	0.0006	0.0009		24	0.0446	0.0433	0.0321	0.0127	0.0331	0.0306	0.0133	0.0154	0.0376	0.0353	0.0409	0.0377	0.0313	0.0397	0.0429	0.0405	0.0323	0.0239	0.0029	0.0291	0.0200	0.0492	0.0486	0.0465	0.0558	0.0574	0.0377
0.8	5	0.0005	0.0004	0.0013	0.0009	0.0012	0.0021	0.0018	0.0014	0.0015	0100.0	0100.0	0.0018	0.0022	0.0022	0.0023	0.0055	070000	0700.0	0.0031	0.0042	0.0050	0.0039	700000	0.0026	0.0026	0.0022	0.0021	0.0009	0.0023		23	0.0160	0.0153	0.0105	0.0025	0.0097	0.0083	0.0010	0.0031	0.0100	0 0095	0.0111	0.0096	0.0071	0.0090	0.0093	0.0084	0.0054	0.0024	0.0068	0.0182	0.0014	0.0102	0.0101	0.0103	0.0130	0.0153	0.0102
ays 7.5	2	0.0002	2000.0	0.0013	0.0011	0.0013	0.0018	0.0017	0.0016	2100.0	1100.0	8100.0	0.0019	0.0022	0.0022	0.0023	10004	10000	0700.0	0.0028	0.0035	0.0039	0,000	E700.0	0.0027	0.0027	0.0023	0.0023	0.0013	0.0022	ys	22	0.0001	0.0023	0.0023	0.0023	0.0034	0.0033	0.0032	0.0046	0.0050	0.0059	0.0067	0.0065	0.0071	0.0077	0.0080	0.0079	0.0073	0.0064	0.0072	0.0072	0.0092	0.0110	0.0110	0.0103	0.0106	0.0091	0.0072
t - t in ds 7.0		0.0003	0.0002	0.0007	0.0005	0.0006	0.0011	0.0010	0.0008	0.0000	0.000	0.0009	0.0010	0.0013	0.0013	0.0013	0.0014	21000	ernn'n	0.0016	0.0024	0.0030	0.0010	2100.0	0.0016	0.0016	0.0014	0.0013	0.0007	0.0014	-t in da	21	0.0001	0.0022	0.0022	0.0020	0.0029	0.0028	0.0025	0.0038	0.0047	0.0047	0.0052	0.0050	0.0054	0.0057	0.0059	0.0058	0.0053	0.0046	0.0053	0.0056	0.0068	0.0079	0.0078	0.0073	0.0074	0.0063	0.0053
ervations 6.5	2	0.0026	0.0025	0.0021	0.0030	0.0029	0.0025	0.0036	0.0046	0.0045	0.0010	0.00.0	0.0048	0.0050	0.0053	0.0055	0.0054	100000	0.0049	0.0043	0.0045	0.0044	0.0050	200000	0.00.0	0.0070	0.0065	0.0067	0.0057	0.0049	edictions	20	0.0002	0.0023	0.0019	0.0013	0.0023	0.0021	0.0014	0.0026	0.0037	0.0036	0.0041	0.0038	0.0039	0.0043	0.0044	0.0043	0.0037	0.0030	0.0029	0.0025	0.0047	0.0060	0.0060	0.0057	0.0059	0.0054	0.0040
SD of obs 6.0		0.0003	0.0003	0.0003	0.0004	0.0004	0.0004	0.0006	0.0007	0.0007	0.0000	0.0008	0.0007	0.0008	0.0008	0.0008	0.0008	200000	0.000	0.0006	0.0009	0.0011	0.0011	1100.0	0.0012	0.0012	0.0011	0.0011	0.0009	0.0008	SD of pro	19	0.000	0.0010	0.0009	0.0007	0.0011	0.0011	0.0009	0.0012	0.0018	0.0018	0.0020	0.0019	0.0020	0.0022	0.0023	0.0023	0.0021	0.0018	0.0017	0.0012	0.0024	0.0031	0.0032	0.0030	0.0031	0.0028	0.0021
RMS 5.5	5	0.0012	0.0014	0.0015	0.0019	0.0019	0.0019	0.0024	0.0029	0.0000	0.0000	0.0032	0.0031	0.0034	0.0036	0.0037	0.0027	00000	00000	0.0033	0.0037	0.0037	6100.0	21-00-0	0.0047	0.0047	0.0043	0.0044	0.0035	0.0033	RM	18	0.0191	0.0116	0.0082	0.0024	0.0081	0.0070	0.0015	0.0041	0.0104	0.004	0.0110	0.0098	0.0083	0.0100	0.0104	0.0097	0.0071	0.0043	0.0022	0.0103	0.0055	0.0131	0.0131	0.0130	0.0151	0.0163	0.0100
5.0		0.0011	0.0005	0.0001	0.0004	0.0002	0.0005	0.0001	0.0005	0.0004	1.0004	GUUU.U	0.0004	0.0002	0.0003	0.0002	0.0001	100000	7000.0	0.0006	0.0011	0.0016	6000.0	700000	0.0004	0.0004	0.0005	0.0006	0.0010	0.0006		17	0.0276	0.0364	0.0268	0.0102	0.0276	0.0252	0.0102	0.0130	0.0322	0.0299	0.0348	0.0319	7920.0	0.0336	0.0360	0.0340	0.0268	0.0195	0.0013	0.0258	0.0173	0.0421	0.0418	0.0402	0.0481	0.0497	0.0320
4.5	2	0.0019	0.0014	0.0006	0.0015	0.0013	0.0005	0.0014	0.0023	0.0001	1700.0	0.0020	0.0022	0.0021	0.0024	0.0024	0.0093	0700.0	0100.0	0.0012	0.0007	0.0000	0.009	7700.0	0.0034	0.0034	0.0032	0.0035	0.0034	0.0022		16	1.0064	0.0255	0.0184	0.0062	0.0184	0.0164	0.0054	0.0083	0.0219	0.0199	0.0232	0.0209	0.0174	0.0216	0.0228	0.0213	0.0163	0.0111	0.0026	0.0215	0.0102	0.0269	0.0268	0.0261	0.0312	0.0332	0.0212
4.0	2	0.0017	0100.0	0.0002	0.0010	0.0008	0.0000	0.0006	0.0014	0.0013	2100 0	G100.0	0.0013	0.0011	0.0013	0.0013	0.0019	710000	0.000	0.0003	0.0005	0.0014	0,0008	000000	0.0018	0.0018	0.0018	0.0020	0.0023	0.0013		15	0.0057	0.0053	0.0032	0.0003	0.0022	0.0015	0.0017	0.000	0.0016	0.0009	0.0011	0.0005	0.0008	0.0005	0.0006	0.0010	0.0020	0.0030	0.0072	0.0114	0.0042	0.0018	0.0018	0.0013	0.0005	0.0014	0.0035
3.5	;	0.0008	0.0006	0.0002	0.0006	0.0005	0.0001	0.0003	0.0008	0.0007	0.000	0.0008	0.0007	0.0007	0.0008	0.0008	0.0008	00000	0.000	0.0004	0.0001	0.0007	0.0005	0100.0	0.0010	0.0010	0.0010	0.0011	0.0012	0.0007		14	0.0020	0.0031	0.0029	0.0024	0.0038	0.0035	0.0029	0.0047	0.0061	10000	0.0069	0.0066	0.0069	0.0075	0.0077	0.0076	0.0068	0.0058	0.0062	0.0061	0.0086	0.0105	0.0105	0.0098	0.0101	0.0089	0.0070
3.0		0.0027	0.0017	0.0004	0.0017	0.0014	0.0000	0.0011	0.0024	0.0099	7700.0	0.0020	0.0023	0.0020	0.0023	0.0023	1600.0	1700.0	4700.0	0.0006	0.0006	0.0019	0.0016	010000	0.0032	0.0032	0.0032	0.0035	0.0039	0.0023		13	0.0061	1000.0	0.0054	0.0041	0.0063	0.0058	0.0044	0.0071	0 0091	0.0000	0 0099	0.0093	0 0095	0.0101	0.0103	0.0101	0.0089	0.0074	0.0074	0.0070	0.0107	0.0132	0.0131	0.0123	0.0127	0.0112	0.0093
		0 6	540	360	480	009	720	840	040	1080	0001	1200	1320	1440	1560	1680	1800	0001	1920	2040	2160	2280	9400	0010	02.92	2640	2760	2880	3000			12	0.0030	0.0033	0.0033	0.0033	0.0042	0.0041	0.0041	0.0052	0.0060	0.0060	0.0066	0.0064	0.0069	0.0072	0.0074	0.0074	0.0071	0.0066	0.0072	0.0074	0.0082	0.0091	0.0091	0.0083	0.0085	0.0067	0.0066
												S.	191	lən	ιu	ti a	C															11	0.000	0.0001	0.0081	0.0063	0.0094	0.0087	0.0068	0.0107	0.0136	0.0134	0.0148	0.0140	0.0143	0.0151	0.0154	0.0151	0.0134	0.0113	0.0116	0.0111	0.0161	0.0196	0.0195	0.0182	0.0188	0.0164	0.0139
																																1		120	240	360	480	009	720	840	090	1080	1200	1320	1440	1560	1680	1800	1920	2040	2160	2280	2400	2520	2640	2760	2880	3000	
																																												SI	ətə	ωı	ų a	;											

Table 13: Section 5.2. For each value of x in the first column, predictions for t > 10 days use data observed at $t \in \{3.5, 4, 4.5, \ldots, 10\}$ days. Observed data correspond to z(x, t) and the prediction is given by the *best fitting quadratic polynomial (solution of a linear least squares problem)*. Each cell of the table shows $|z_{\text{pred}}(x,t) - z(x,t)|$. In the left-hand part of the table, values of z(x,t) correspond to synthetic trainind data (used in the fitting), while in the right-hand part of the table z(x,t) correspond to synthetic (testing) data that is not being used in the fitting. The last line in the table shows the RMSD for each t.

RMISD of observation 3.0 3.5 4.0 4.5 5.0 5.5 6.0 6.5 120 0.0004 0.0003 0.0004 0.0001 0.0001 0.0011 0.0011 120 0.0004 0.0003 0.0004 0.0002 0.0001 0.0011 0.0011 240 0.0002 0.0002 0.0002 0.0002 0.0012 0.0011 0.0013 0.0011 360 0.0002 0.0002 0.0001 0.0002
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3.0 3.5 4.0 4.5 5.0 120 0.0004 0.0004 0.0003 0.0004 0.0005 0.0001 0.0005 0.0001 0.0005 0.0001 0.0005 0.0005 0.0005 0.0001 0.0005 0.0001
3.0 3.5 4.0 4.5 3.0 3.5 4.0 4.5 3.1 3.5 4.0 4.5 3.1 3.0 0.0014 0.0003 0.0001 3.60 0.0004 0.0003 0.0002 0.000 3.60 0.0004 0.0003 0.0002 0.000 3.60 0.0003 0.0002 0.0001 0.000 3.60 0.0005 0.0004 0.0003 0.001 440 0.0005 0.0004 0.0003 0.001 960 0.0005 0.0004 0.0003 0.001 960 0.0005 0.0004 0.0003 0.001 960 0.0005 0.0004 0.0003 0.001 960 0.0005 0.0004 0.0003 0.001 1580 0.0006 0.0004 0.0003 0.001 1580 0.0006 0.0006 0.0003 0.001 1580 0.0006 0.0006 0.0003
3.0 3.5 4.0 120 0.0004 0.0004 0.000 2400 0.0004 0.0003 0.000 360 0.0004 0.0004 0.000 360 0.0004 0.0004 0.000 360 0.0004 0.0003 0.000 360 0.0004 0.0004 0.000 360 0.0004 0.0004 0.000 360 0.0006 0.0004 0.000 360 0.0006 0.0004 0.000 360 0.0006 0.0004 0.000 360 0.0006 0.0004 0.000 360 0.0006 0.0004 0.000 360 0.0006 0.0004 0.000 3720 0.0006 0.0004 0.000 3180 0.0006 0.0004 0.000 3180 0.0006 0.0006 0.000 3240 0.0008 0.0006 0.000 3240 0.0006 0.0
10 0.0001 0.0003 3.0 3.5 120 0.0004 0.0003 0.0004 0.000 240 0.0004 0.0003 0.0003 0.0003 360 0.0004 0.0003 0.0003 0.0003 360 0.0004 0.0003 0.0003 0.0003 720 0.0005 0.0006 0.0006 0.0003 960 0.0006 0.0006 0.0006 0.0006 11280 0.0006 0.0006 0.0006 0.0006 11280 0.0006 0.0003 0.0006 0.0006 2240 0.0006 0.0003 0.0006 0.0006 2240 0.0006 0.0003 0.0003 0.0003 2240 0.0003 0.0003 0.0003 0.0003 2240 0.0003 0.0003 0.0003 0.0003 2240 0.0003 0.0003 0.0003 0.0003 2240 0.0003 0.0003 0.0003 0.0003
3.0 3.0 3.1 0.0002 3.1 0.0002 3.1 0.0002 3.1 0.0002 3.1 0.0002 3.1 0.0002 3.1 0.0002
0 0 1200 360 2400 360 2800 480 6900 480 7200 720 11260 720 11260 1200 11260 22640 11260 1200 11260 1220 111 12 1200 1000 11 12 11 12 11 12 11 12 11 12 11 12 11 12 11 12 11 12 11 12 11 12 11 12 11 12 11 12 11000 0.0005 11000 0.0005 11000 0.0005 11200 0.0005 11200 0.0005 11200 0.0005
11 11 11 11 11 11 110 0.0014 110 0.0056 110 0.0057 110 0.0057 110 0.0057 110 0.0057 110 0.0057 110 0.0057 1100 0.0057 1120 0.0057 1120 0.0057 1120 0.0057 1120 0.0057 1120 0.0057 1120 0.0057 1120 0.0057 1120 0.0057 1120 0.0057 1120 0.0057 1120 0.0057 1120 0.0057 1120 0.0057 1120 0.0057 1120 0.0057 1120 0.0057
0 120 1200 1200 1200 11200 11200 11200 11200

Table 14: Section 5.2. For each value of x in the first column, predictions for t > 10 days use data observed at $t \in \{3.5, 4, 4.5, \ldots, 10\}$ days. Observed data correspond to z(x, t) and the prediction is given by the *best fitting cubic polynomial (solution of a linear least squares problem)*. Each cell of the table shows $|z_{pred}(x,t) - z(x,t)|$. In the left-hand part of the table, values of z(x,t) correspond to synthetic trainind data (used in the fitting), while in the right-hand part of the table z(x,t) correspond to synthetic (testing) data that is not being used in the fitting. The last line in the table shows the RMSD for each t.